



Operational forecast system for storm surges in Brazil and Mozambique

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Southwest Indian Ocean



- Madagascar and Mozambique Channel
- wide and shallow continental shelf (central and south)
- large astronomical tidal range
- large rivers (central and south)
- immense low lying coastal plains
- high vulnerability to tropical cyclones

Tropical cyclones in Southwest Indian Ocean between 1980-2007 (November - April) and climatological SST (Mavume et al. 2009)









Why to model storm surges in the central coast of Mozambique?



Landfalling Tropical cyclones in Mozambique and Madagascar



Scientific questions to be explored:

- Can increased model resolution improve tidal representation and thus total water levels?
- Are there differences in using 3D (baroclinic) or 2D (barotropic) versions of the model to simulate storm surges in this specific area?
- Are there combined effects of the forcings: tides, winds and sea level pressure?

Modelling Strategy

The Princeton Ocean Model on the SouthWest Indian Ocean (POM-SWIO)



POM-SWIO grid



Validation – Tides

Amplitudes (cm) and phases (°GW) of the principal semidiurnal tidal constituents



• Both model configurations are able to capture the basin-wide large-scale tidal patterns;

• But, only when the horizontal resolution is as high as 1/12° (~8 km) the model is able to capture tidal amplification in central part of Mozambique coast

Validation – Tides

Principal tidal constituents bias (model minus observation)

	M2		S2		01		К1	
	Amplitude	Phase	Amplitude	Phase	Amplitude	Phase	Amplitude	Phase
	(m)	(°)	(m)	(°)	(m)	(°)	(m)	(°)
	Observation							
	1.8184	28.55	1.0317	138.49	0.0223	287.38	0.0074	298.70
	Model bias							
3D 1/6°	-0.35 (19%)	-16.81	-0.18 (17%)	-19.88	0.01 (0.44%)	2.37	<-0.01 (57%)	-185.22
2D 1/6°	-0.33 (18%)	-16.05	-0.16 (16%)	-19.21	0.01 (0.46%)	1.44	<-0.01 (66%)	-178.13
3D 1/12°	0.13 (7%)	-4.73	0.14 (14%)	-6.19	0.01 (59%)	6.59	<-0.01 (64%)	-154.91
2D 1/12°	0.15 (9%)	-0.78	0.15 (14%)	-0.57	0.01 (55%)	9.40	<-0.01 (59%)	-149.71



Validation – Surges

TC Lisette (February/March 1997)



00

1.0

0.02

$$MAE = \frac{1}{n} \sum_{i=1}^{n} |M_i - O_i| \qquad Skill = 1 - \frac{\sum_{i=1}^{n} (M_i - O_i)^2}{\sum_{i=1}^{n} (M_i - \overline{O}| + |O_i - \overline{O}|)^2}$$

Model performance for previous TC cases within the SWIO domain

(when data exists for comparisons)

Tropical cyclone Lisette (1997)



Tropical cyclone Bonita (1996)

Tropical Cyclone Daniella (1996)



A recent storm surge case:

The catastrophic tropical cyclone Idai: 04 - 16 March 2019

TC Idai (March 2019)

Most disastrous event in Mozambique's recent past:

- Intense TC (63 m/s; cat 3 on Saffir-Simpson hurricane scale);
- Destroyed over 80% of the second major city of Mozambique (Beira);
- More than 1 million people affected;
- Over 600 deaths reported;
- Widespread flood lasted for more than a week.











TC Idai (March 2019)



The influence of tides and SLP on surges

Factor separation [Stein and Alpert (1993) methodology]

Experiment	Forcing
expA (control)	wind + tides + SLP
expB (no SLP)	wind+ tides
expC (no Tide)	wind+ SLP
expD	wind

- **r1 = expB expD** residual tidal due to tide-surge interaction
- r2 = expC expD residual due to SLP
- **r12 = expA (expB+expC) + expD** residual due to tide-surge-SLP interaction

The influence of tides and SLP on surges peak – Lisette (1997)



• All the terms acted to increase the surge height during the peak;

• SLP term seems to dominate within all the experiments, but it is slight higher in the higher resolution configuration;

The influence of tides and SLP on surges peak – Idai (2019)



Issues to complete the setup for an operational stage

Final adjustment of tidal constituents

Innitial and boundary fields (for external and internal modes) from global analysis and forecasts Use of ensemble forecast to evaluate uncertainties in the

atmospheric forcing fields

POM_global 1 x 1 degree





TPXO 7.2





S2



Preliminar Conclusions

- Modeled (1/12°) distributions of the semidiurnal tidal amplitudes and phases are in very good agreement with the observations as well as with the reference given by TPXO for both 2D and 3D particularly for 1/12° grid application
- The 3D configuration tends to reproduce higher storm surges heights than 2D during storm surge peaks, in part due to tide-surge nonlinear interactions

Preliminar Conclusions

- After the wind, SLP residual term appeared to dominate in the TC Lisette (1997) storm surge case whereas tide-surge nonlinear interactions surpassed SLP residual term in the TC Idai (2019) case. However it is necessary to reinforce that minimum SLP was largely misrepresented by the atmospheric forcing
- Dependence on global forecasts for initial and boundary conditions from Copernicus or others; solution for external mode could come from own global system, but still there is dependence on internal modes (T,S and currents)

Thanks for your attention!

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