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Determination of Oceanic Extremes using a Spatial Ensemble of Satellite Data



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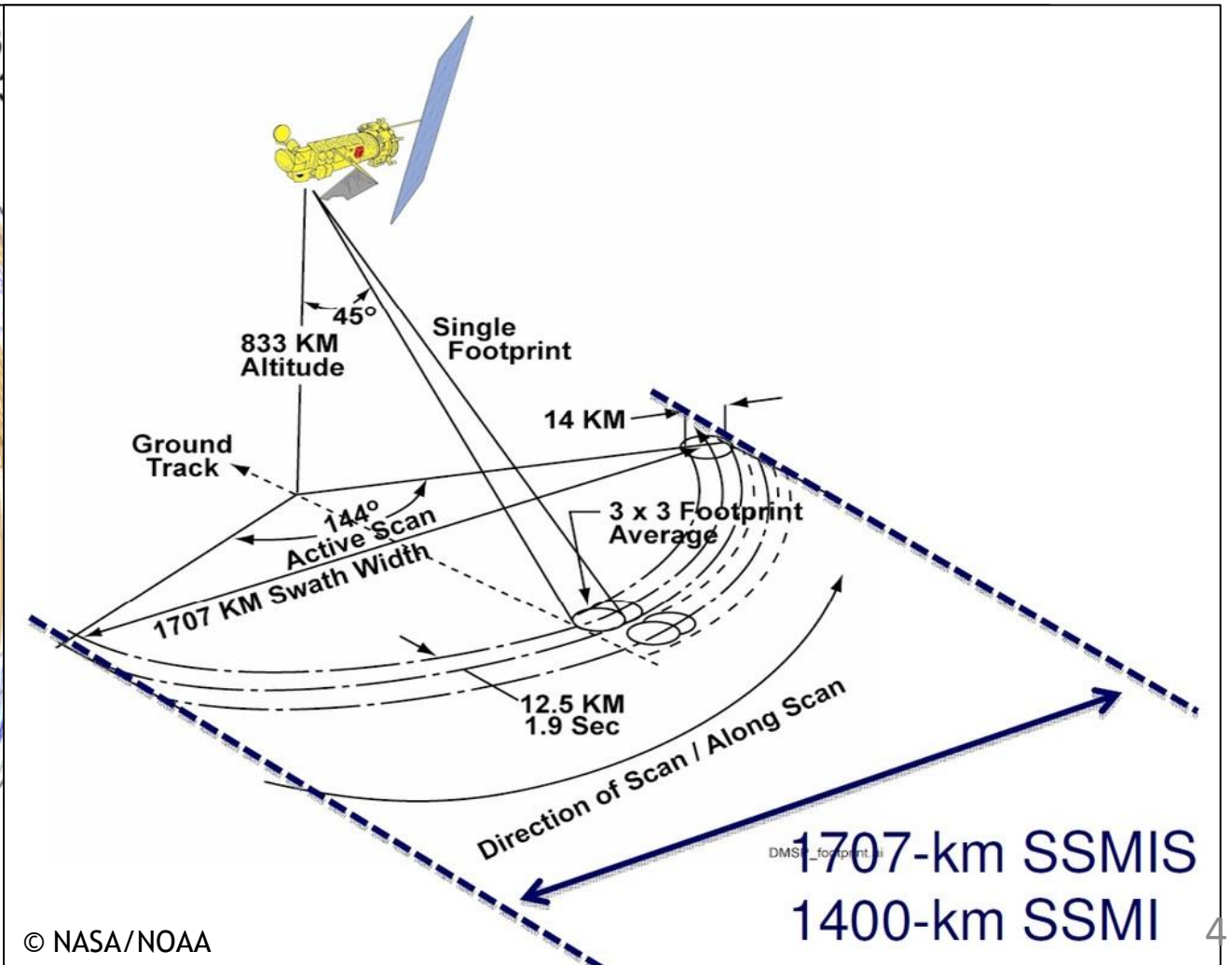
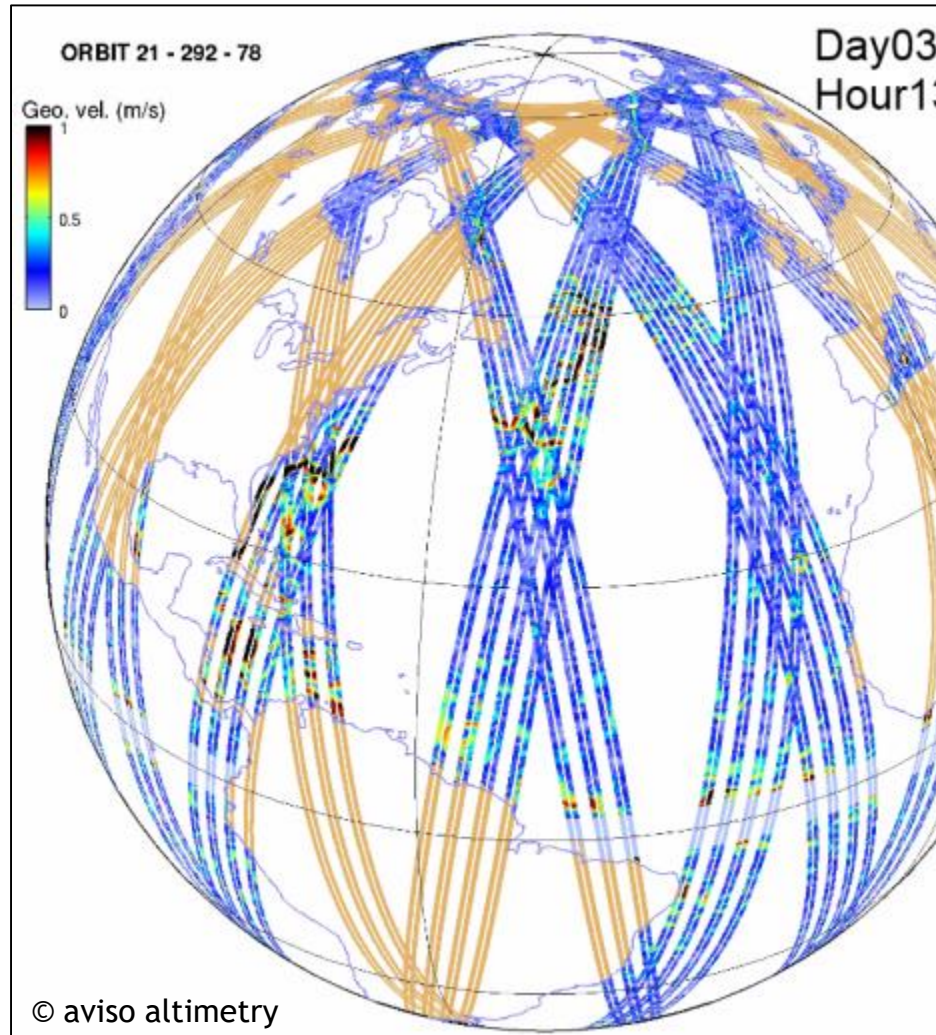
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1 Global distribution of extreme wave height and wind speed from satellite data (Takbash et al. 2019)

- Satellite data (Young et al., 2017):
 - Altimeter/Radiometer → Long duration records & global coverage
- Two general approaches:
 - Initial Distribution Method (IDM) & Peaks over Threshold (PoT)
- Focus:
 - How well does the PDF fits the data?
 - How well is the tail of the PDF defined?

1 Global distribution of extreme wave height and wind speed from satellite data (Takbash et al. 2019)



Correction of the Radiometer data

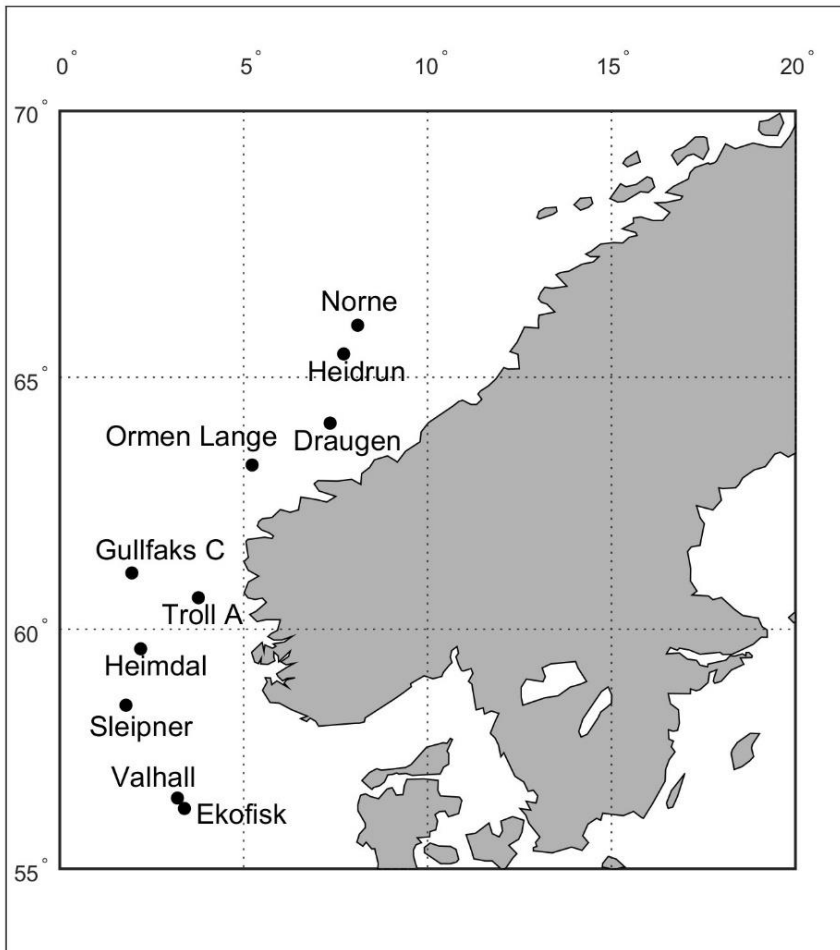


Figure 1 Locations of offshore platforms

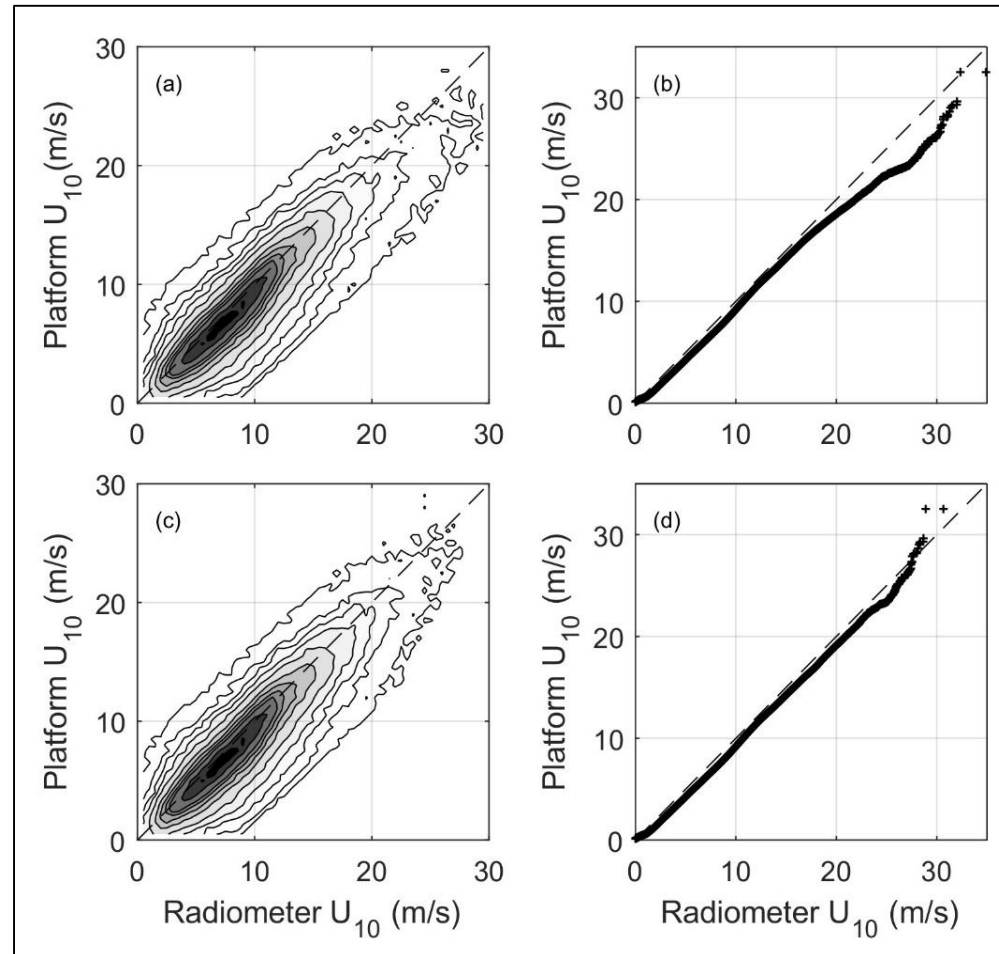
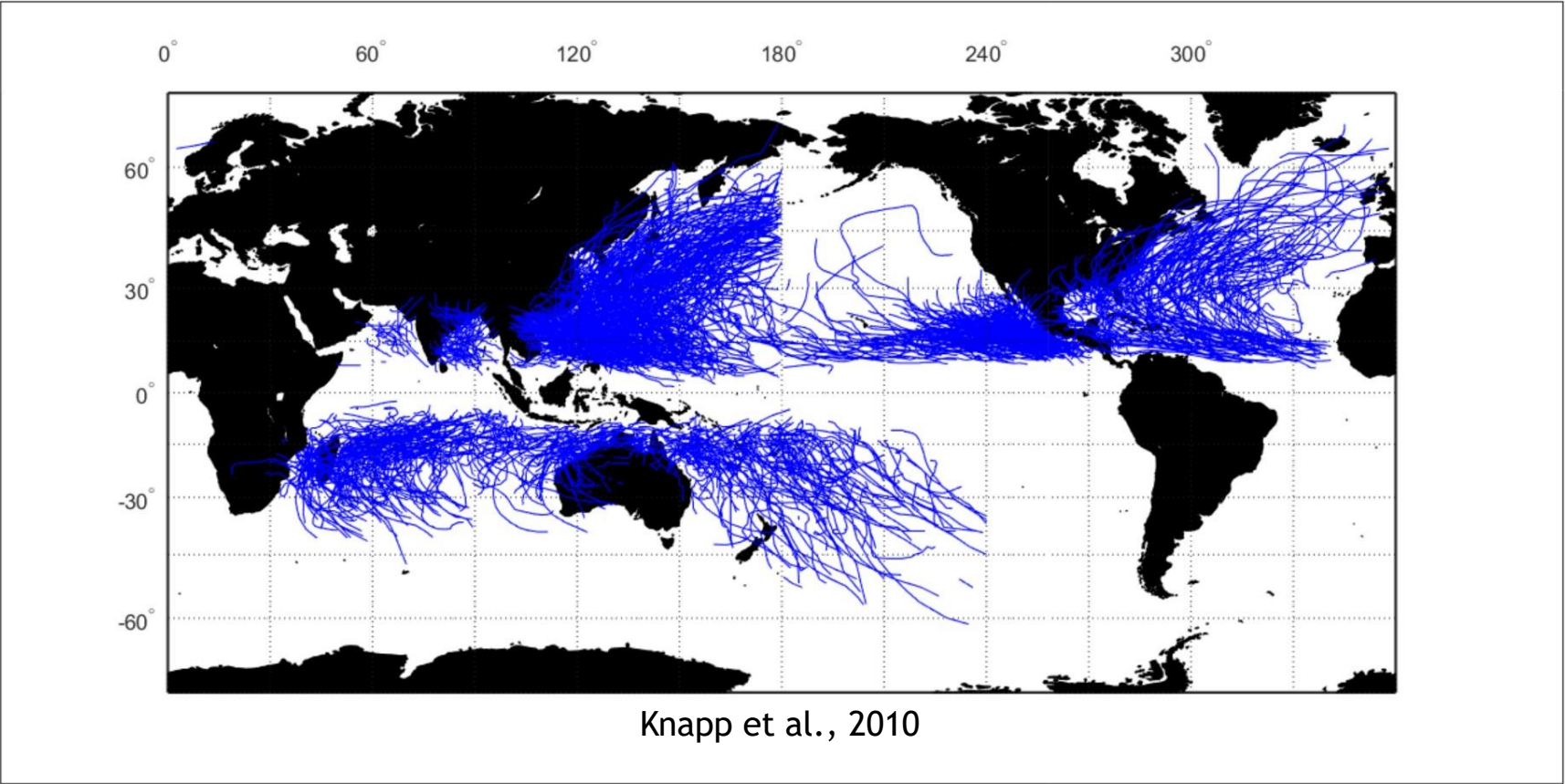


Figure 2 Radiometer-platform comparisons

Global distribution of Extremes - Altimeter PoT

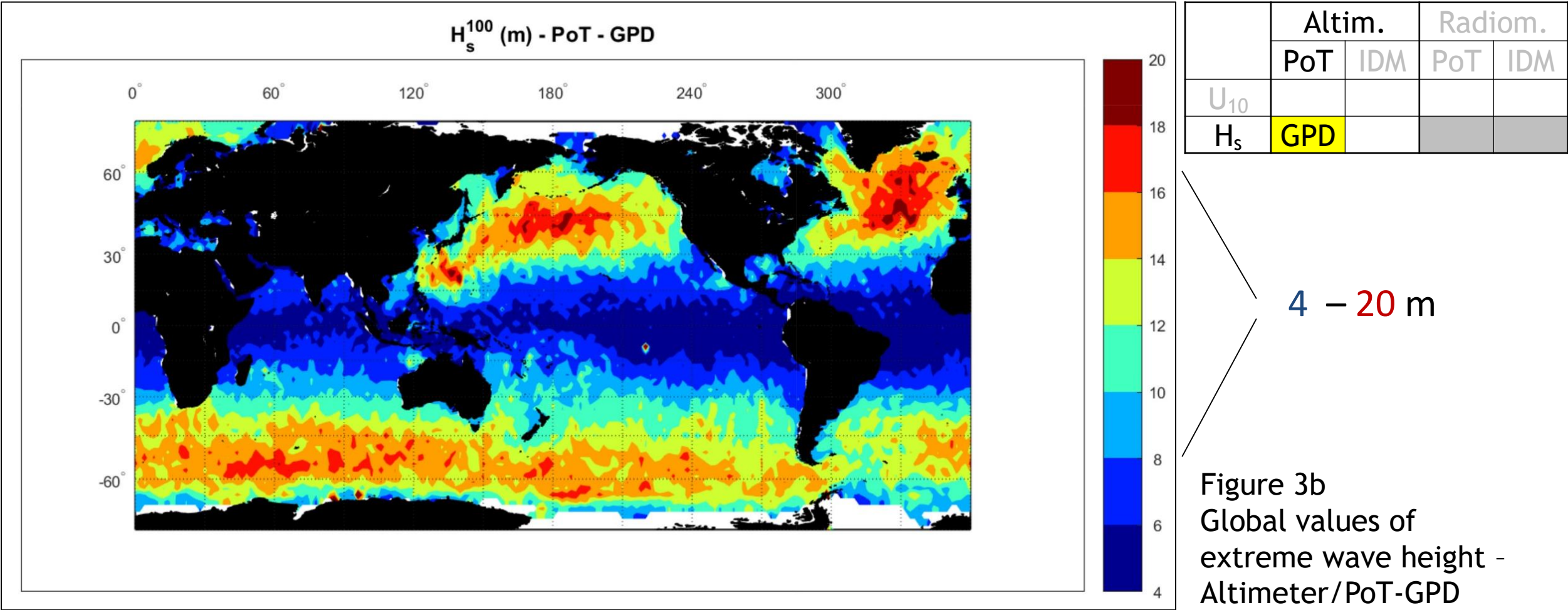


	Altim.		Radiom.	
	PoT	IDM	PoT	IDM
U ₁₀	GPD			
H _s				

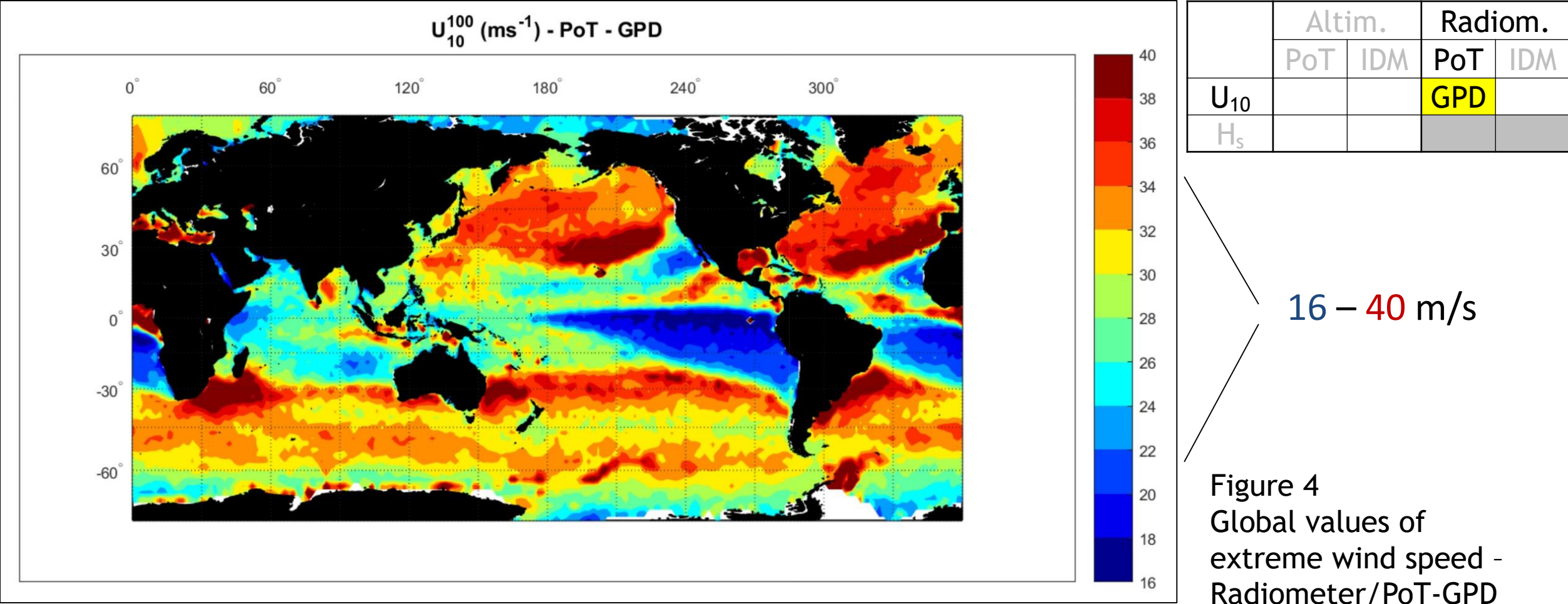
16 – 40 m/s

Figure 3a
Global values of
extreme wind speed -
Altimeter/PoT-GPD

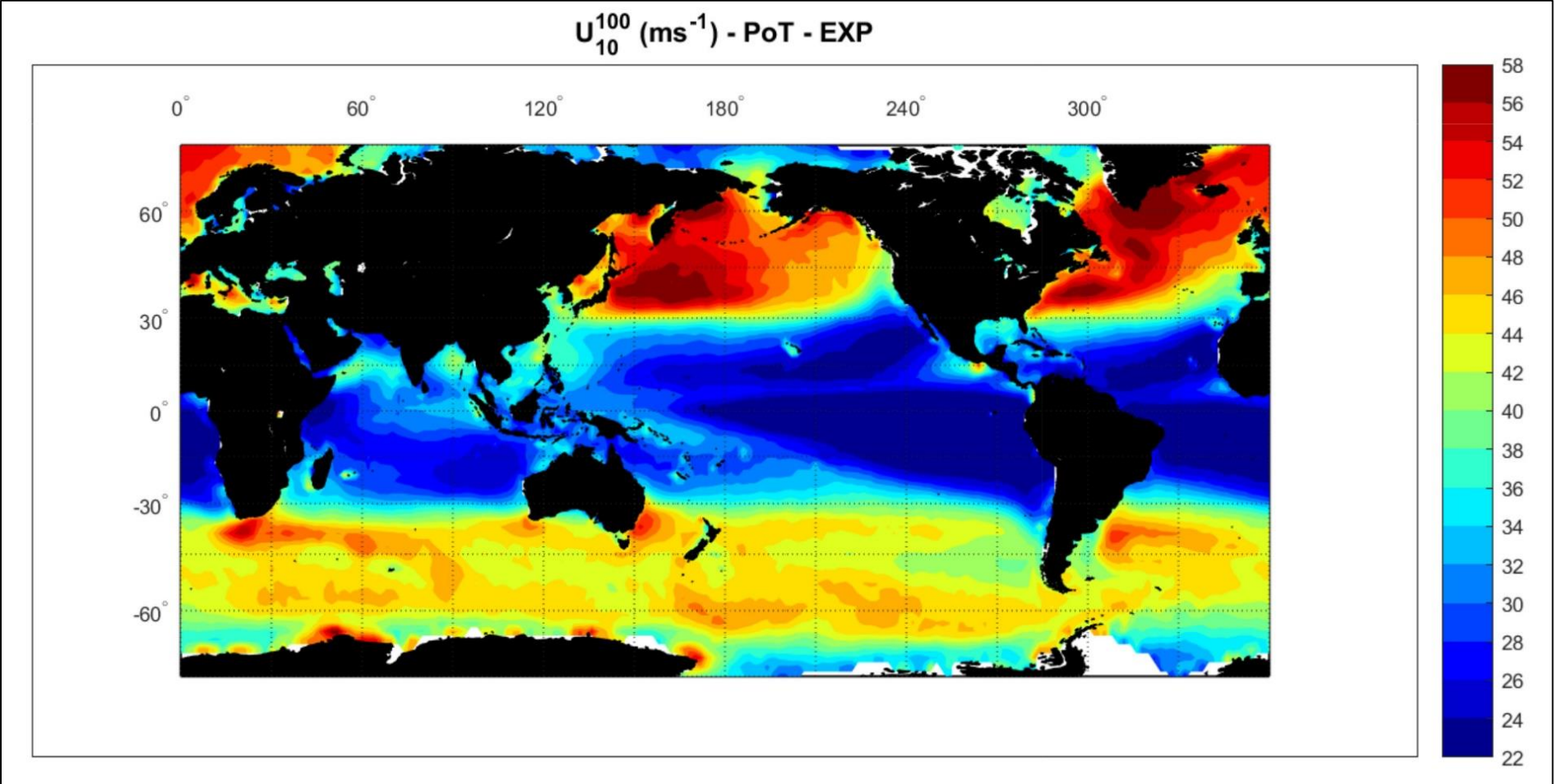
Global distribution of Extremes - Altimeter PoT



Global distribution of Extremes - Radiometer PoT



Global distribution of Extremes - Radiometer PoT

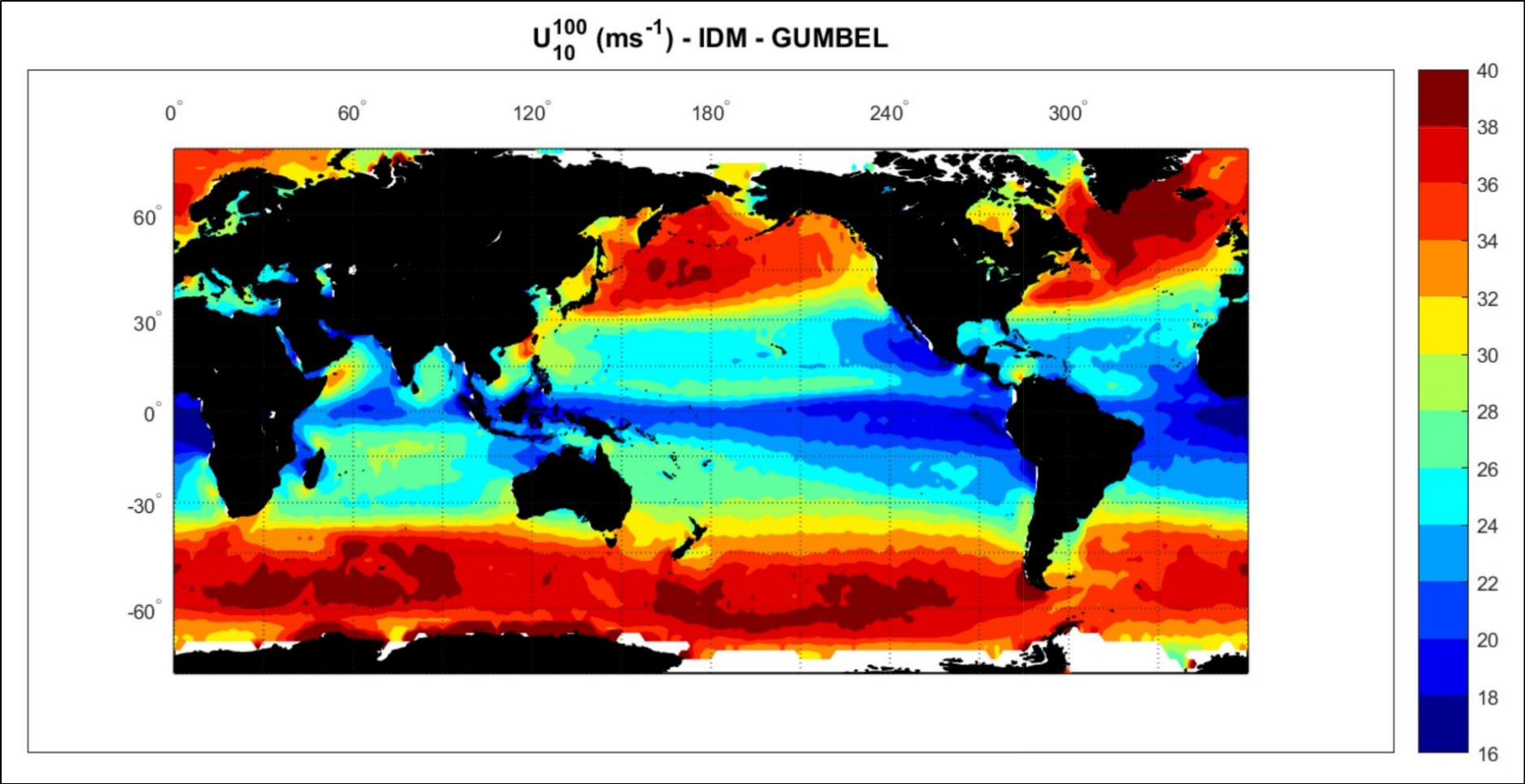


	Altim.		Radiom.	
	PoT	IDM	PoT	IDM
U_{10}			EXP	
H_s				

!! 22 – 58 m/s

Figure 5
Global values of
extreme wind speed -
Radiometer/PoT-EXP

Global distribution of Extremes - Altimeter IDM

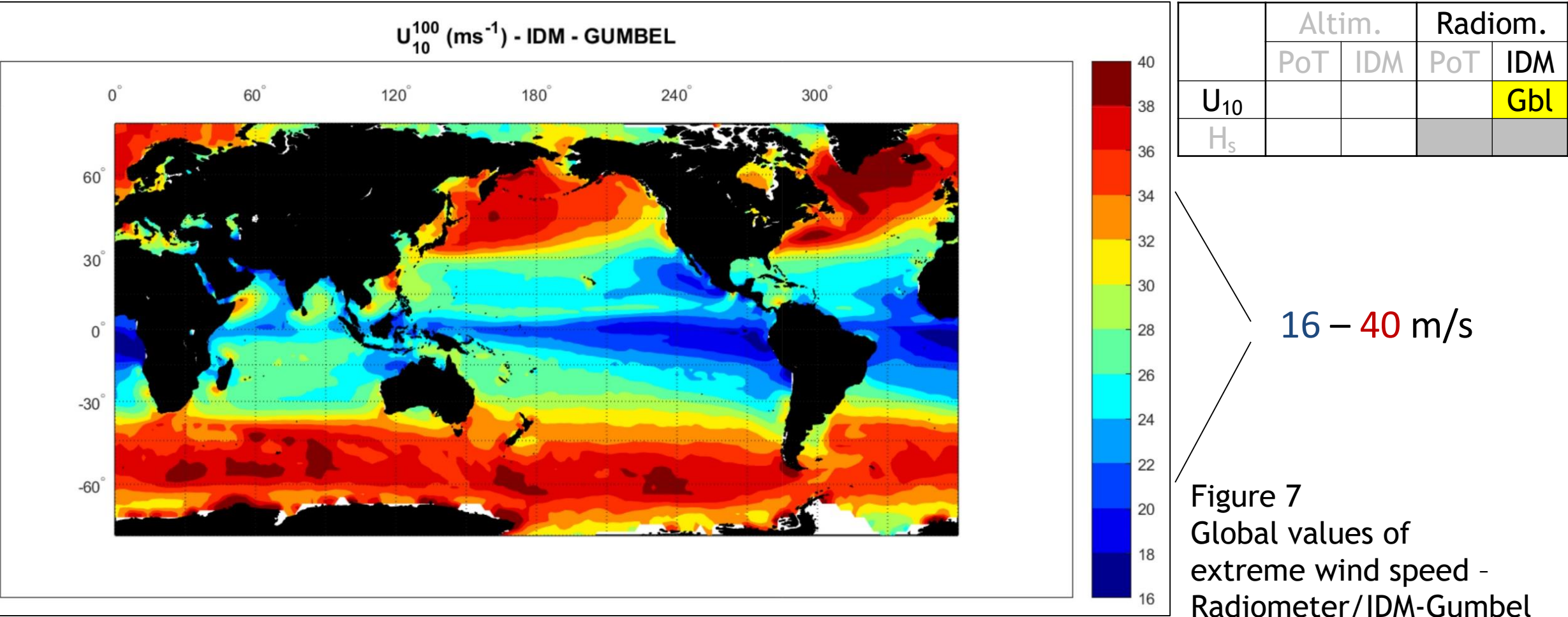


	Altim.		Radiom.	
	PoT	IDM	PoT	IDM
U_{10}		Gbl		
H_s				

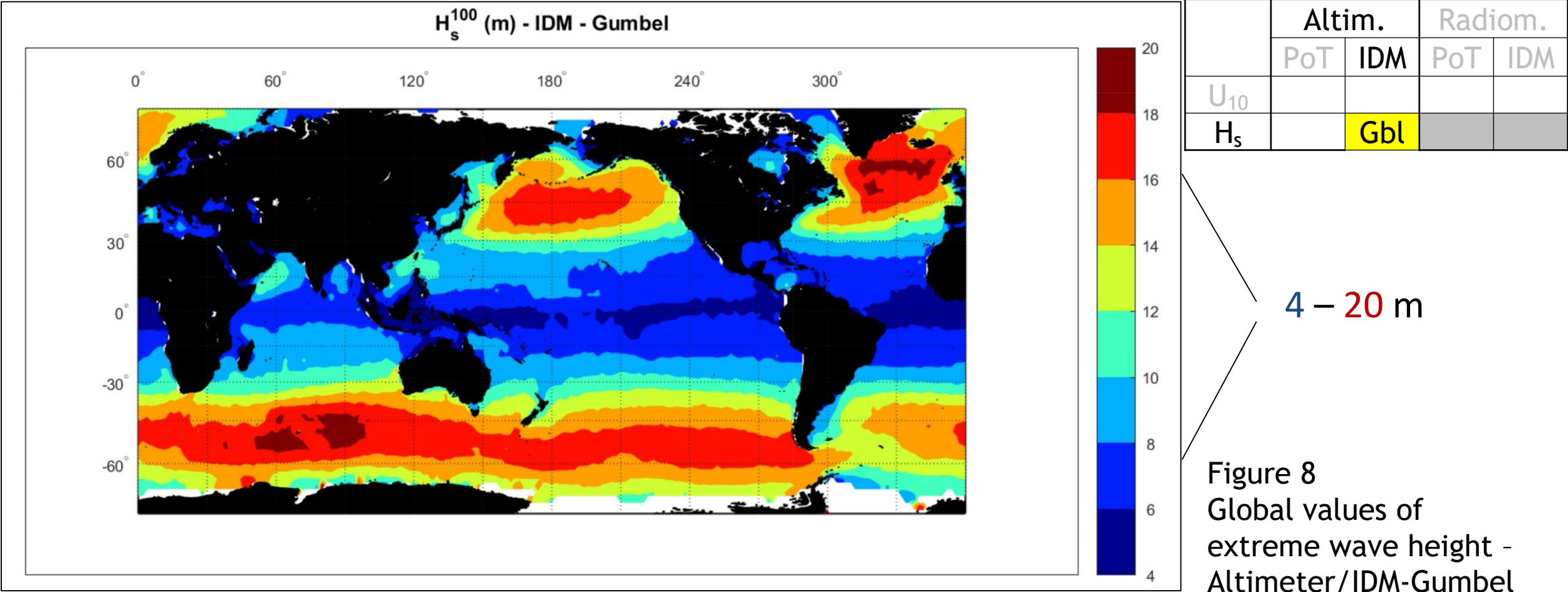
16 – 40 m/s

Figure 6
Global values of extreme
wind speed -
Altimeter/IDM-Gumbel

Global distribution of Extremes - Radiometer IDM



Global distribution of Extremes - Altimeter IDM



2 Conclusion

- ❑ The new Satellite data enables **PoT analysis for the first time for**
 - **Altimeter:**
 - Values consistent with buoy and previous numerical model data
 - Much greater fine scale structure
 - **Radiometer:**
 - Unacceptable “fair-weather” bias
 - **Unusable for PoT**
-
- ❑ IDM yield quite biased estimates of extremes and their spatial distribution
 - **Comparing to PoT, little reason to use IDM in the future!**

3 Global extreme wave height from spatial ensemble data (Takbash and Young 2019)

Aim:

reduce potential errors and **the size of confidence limits** on the resulting estimates of extremes when applying the **PoT approach** to **altimeter data**

→ Potential method: **increase the amount of data points**

by combining data from adjacent regions

→ **spatial ensemble** of aggregated (pooled) data

- Equivalent to a longer duration time series!
- Method has been tested by Breivik et al. (2013; 2014) in the time domain

3 Global extreme wave height from spatial ensemble data (Takbash and Young 2019)

Two criteria must be satisfied:

1. The areas must be far enough apart that the data points of these areas are **independently distributed** (i.e. uncorrelated)
2. The areas must still be representative of a similar wave climate by showing a **common parent distribution**

→ Assessment: ERA-Interim reanalysis data (continues in spatial/time domain)

3 Global extreme wave height from spatial ensemble data (Takbash and Young 2019)

To satisfy the **first criterion**:

Determine the **spatial coherence of wave height** on a global basis using the Pearson correlation coefficient

- Removed the **monthly variation** (monthly mean)
 - Deseasonalised time series
 - Irregular variations (storms and long-term trend)

Correlation Ellipses for the irregular variation of H_s

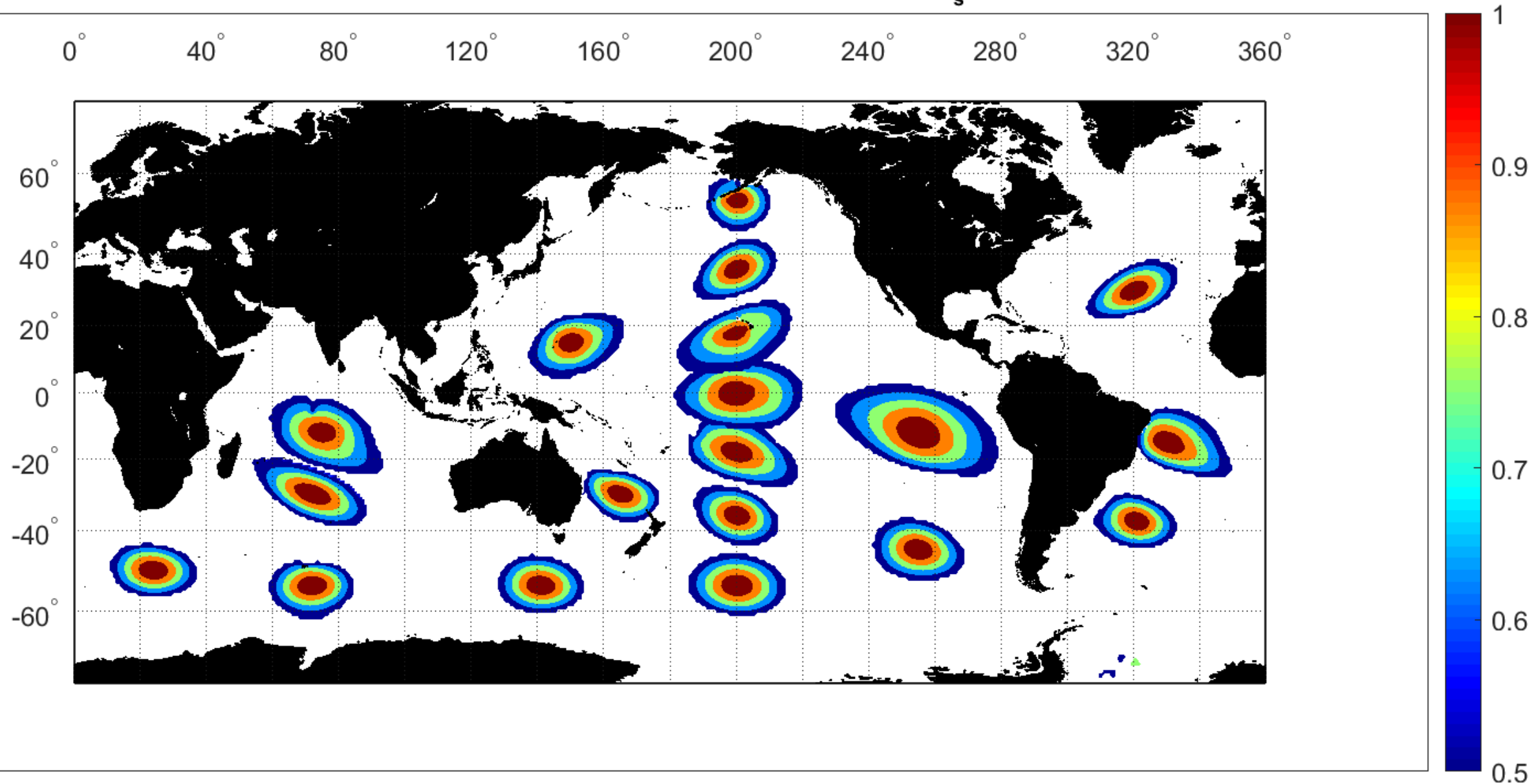


Figure 9 Correlation Ellipses - irregular variation of H_s [monthly means subtracted from the time series]

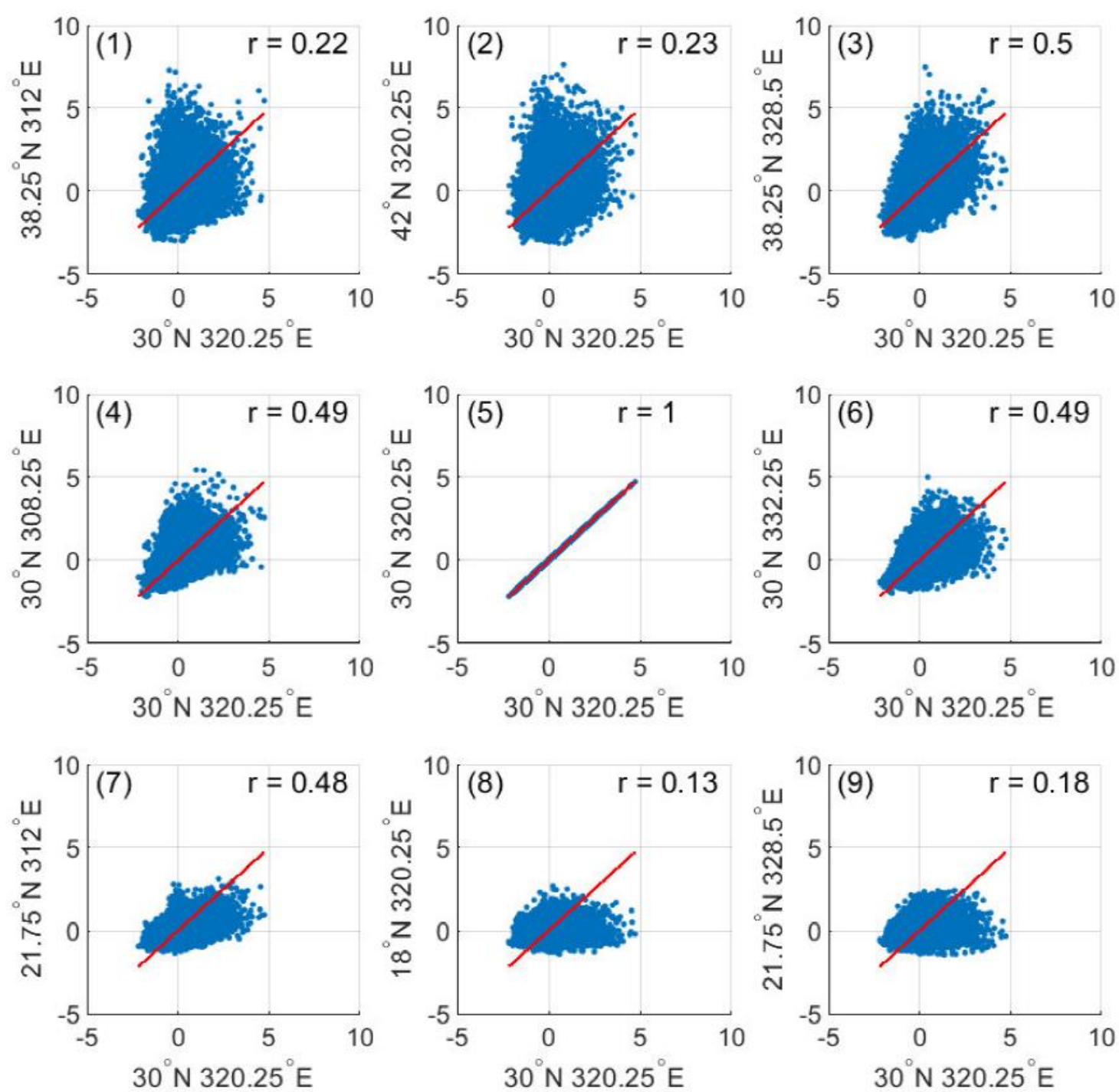


Figure 10 Scatter plots of deseasonalised H_s between a location at 30°N, 320.25°E (North Atlantic) and surrounding locations (12° radius)

3 Global extreme wave height from spatial ensemble data (Takbash and Young 2019)

→ Assessment of spatial coherence between areas in zonal direction!

Second criterion:

The relative percentage difference (RPD) between

- monthly means (averaged over the duration of the records)
- averaged monthly 99th percentile values

→ Wave climate is similar when **RPD** is less than 10 % for both parameter!

$$\overline{\text{RPD}}(i, j) = \frac{1}{12} \sum_{k=1}^{12} \frac{|\overline{H}_s(i, k) - \overline{H}_s(j, k)|}{\overline{H}_s(i, k)},$$
$$\text{RPD}^{99}(i, j) = \frac{1}{12} \sum_{k=1}^{12} \frac{|H_s^{99}(i, k) - H_s^{99}(j, k)|}{H_s^{99}(i, k)}.$$

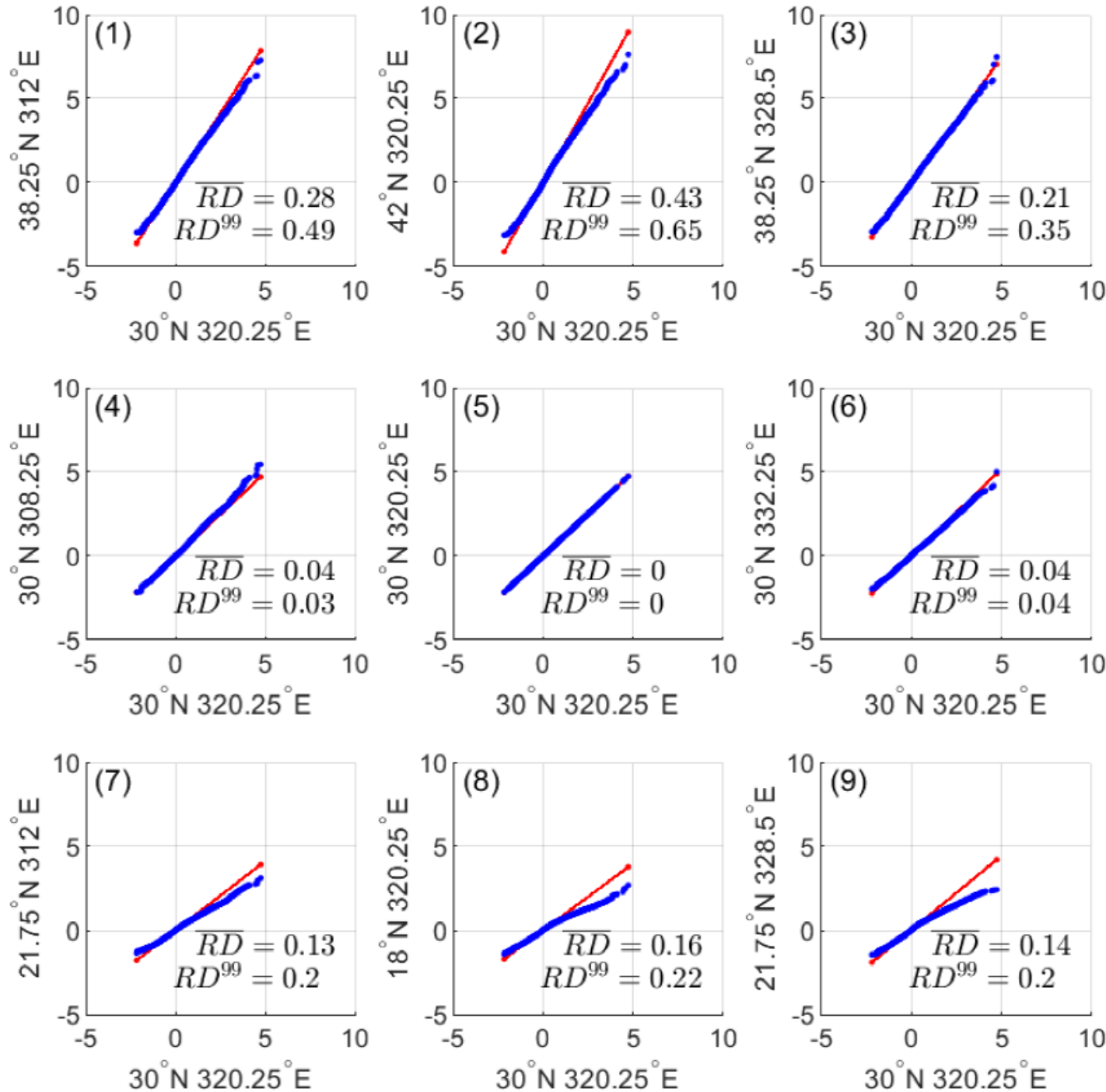


Figure 11 QQ plots of deseasonalised H_s between a location at 30° N, 320.25° E (North Atlantic) to surrounding locations (12° radius)

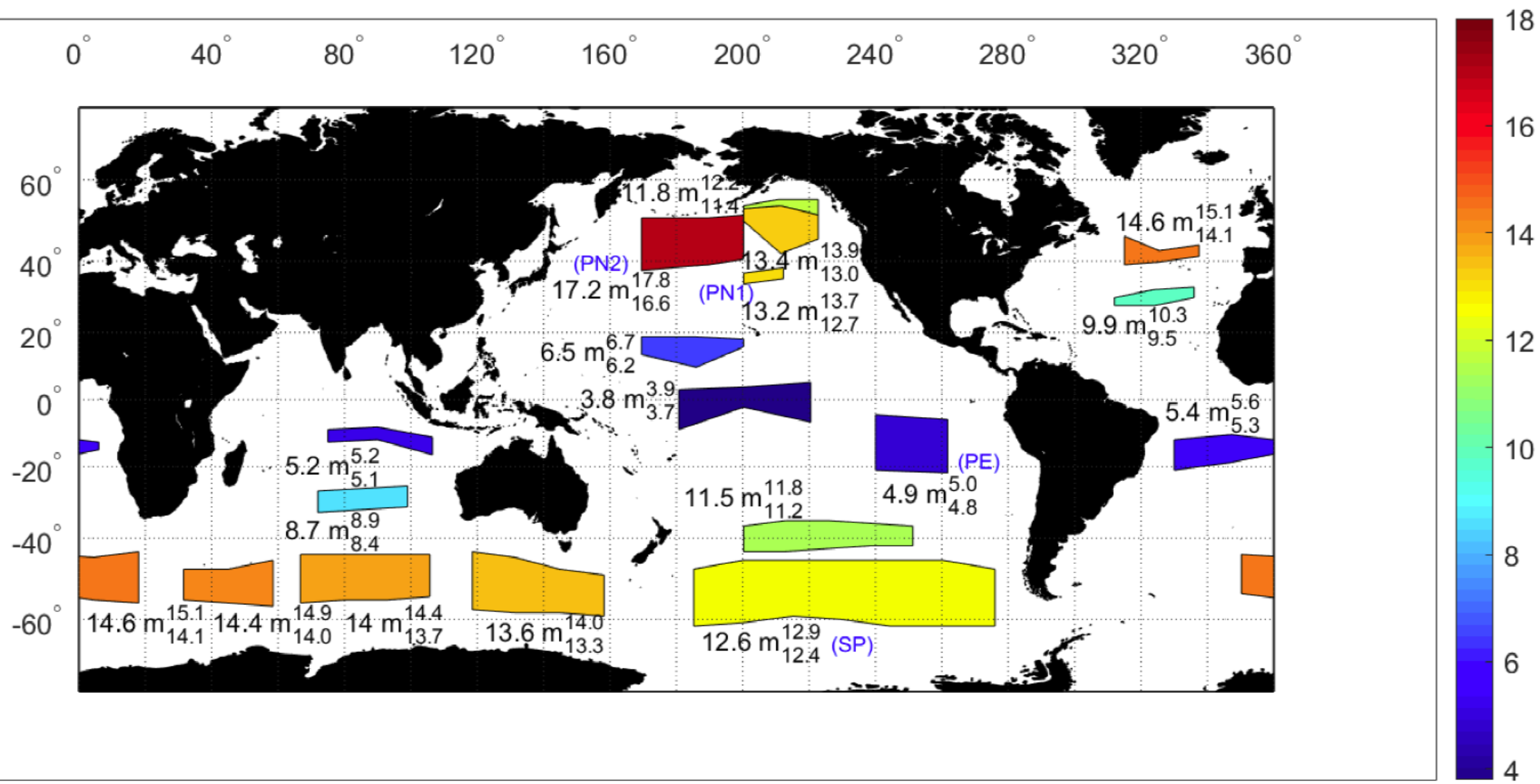


Figure 12 Ensemble spatial regions for ERA-Interim data

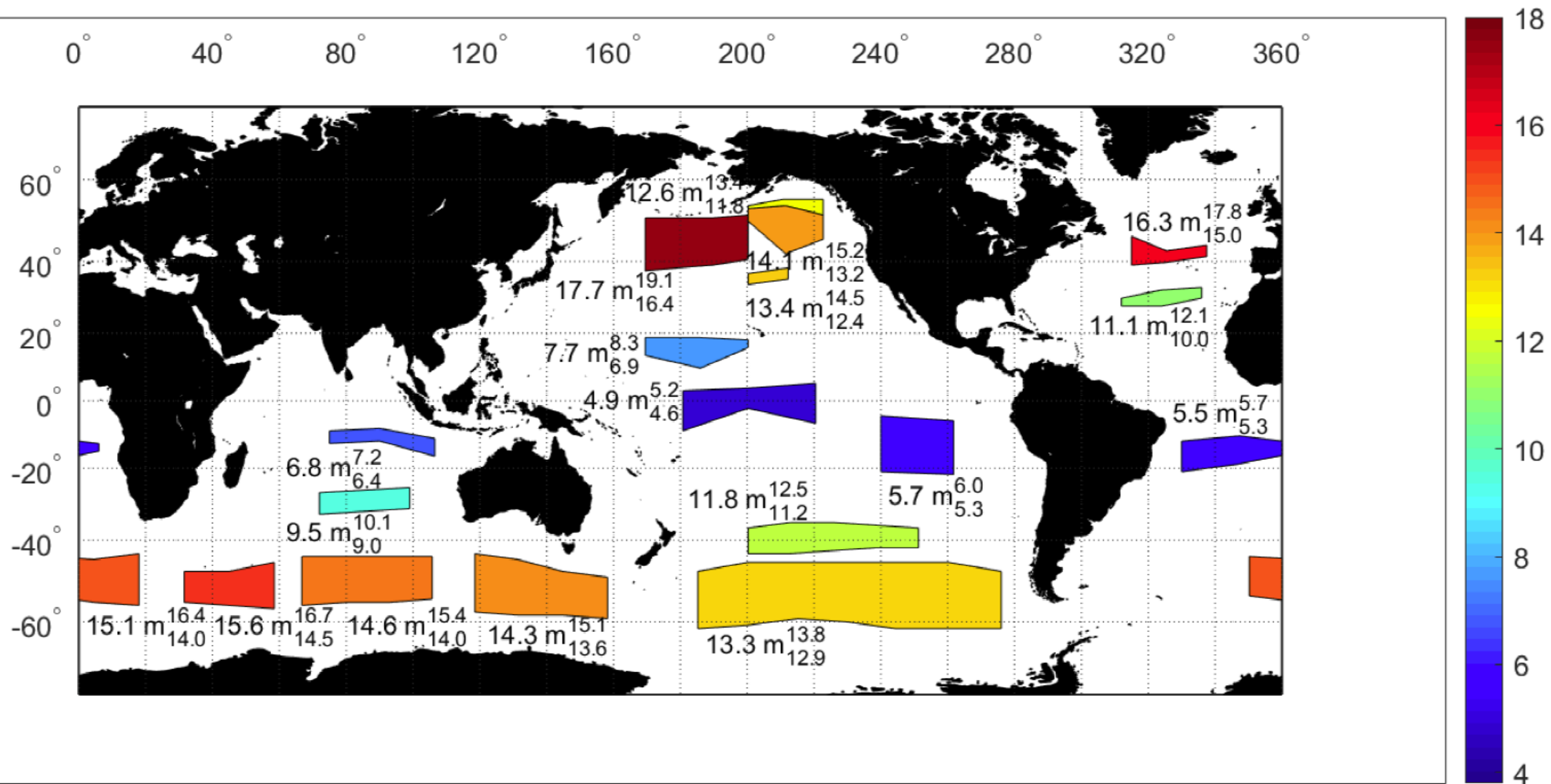


Figure 13 Ensemble spatial regions for **Altimeter** data

4 Conclusion

❑ For **ERA-Interim data** we created spatial ensemble equivalent to:
60 yrs (N Pacific) to 210 yrs (S Pacific)

❑ For **Altimeter data** we created spatial ensemble equivalent to:
54 yrs (N Pacific) to 189 yrs (S Pacific)

→ **Reduced size of the confidence interval by ~30% - 60% depending on the size of the spatial ensemble data**