

Estimation of directional spreading within surface water waves

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Introduction

- Crest height statistics important for design of marine structures;
- Evidence of amplifications beyond 2nd-order from the field; (Christou & Ewans, 2012)
- Directionality acts to reduce these non-linear amplifications;
- Limited amount of data from the field in steep sea-states;
- Laboratory simulations incorporating non-linearity and directionality;
- Role of directionality in steep sea-states and the formation of large wave events.



Key points

1. Generation of directionally spread sea-states in a laboratory environment;
2. Estimation of the degree of directional spreading within a sea-state and large individual waves;
3. Changes in the directional spreading with increasing steepness.



Experimental investigation



Imperial College London
wave basin

Water depth: $d = 1.25$ m

Measurements: η and u, v, w .

20 seeds x 1024 s

JONSWAP spectra, $\gamma = 2.5$

$H_s = [0.10, 0.15, 0.20]$ m

$T_p = 1.6$ s, $\sigma_\theta = 15^\circ$

For the frequency spectrum, $S_{\eta\eta}(\omega)$, the directional spectrum is:

$$F(\omega, \theta) = S_{\eta\eta}(\omega)D(\omega, \theta), \quad (1)$$

where $D(\omega, \theta)$ is the directional spreading function (DSF).

In this study: Gaussian DSF - frequency independent

$$D(\omega, \theta) = \frac{A}{\sigma_\theta \sqrt{2\pi}} e^{\left[-\frac{(\theta - \theta_m)^2}{2\sigma_\theta^2}\right]} \quad (2)$$

with σ_θ the standard deviation, θ_m the mean wave direction and A a scaling parameter.



Methods of directional wave generation: DSM

Three methods of generation:

- Double Summation Method (DSM):
 - Each frequency ($i=1:N$) has M wave components at different directions
 - Total MN components
 - Issues with ergodicity, cancellation of wave components with same frequency but different directions

$$\eta(x, y, t) = \sum_{i=1}^N \sum_{j=1}^M A_{ij} \cos[\omega_i t - k(x \cos \theta_j + y \sin \theta_j) + \epsilon_{ij}] , \quad (3)$$

where

$$\omega_i = i(2\pi\Delta f), \theta_j = j \Delta\theta \text{ and } A_{ij} = \sqrt{2 S_{\eta\eta}(\omega, \theta) \Delta\omega \Delta\theta}$$

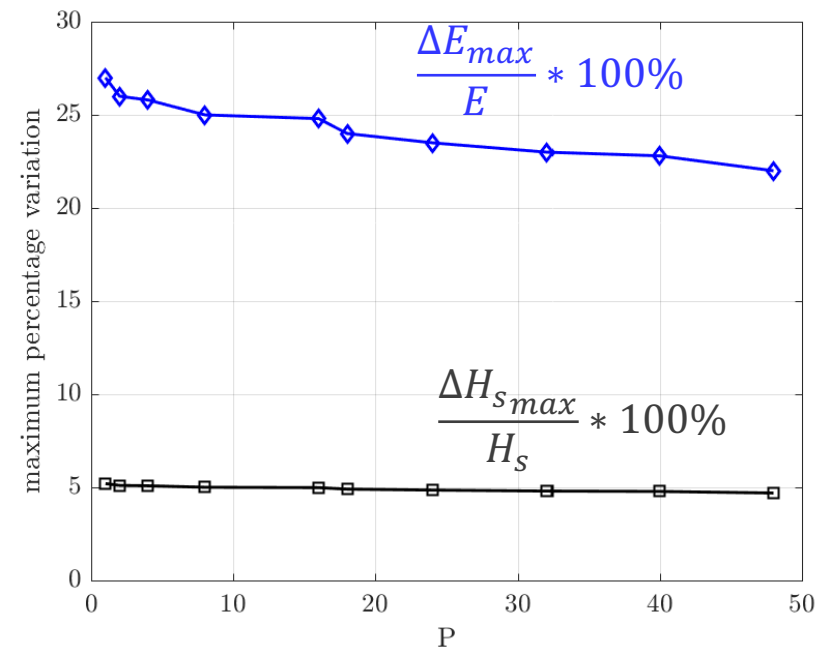
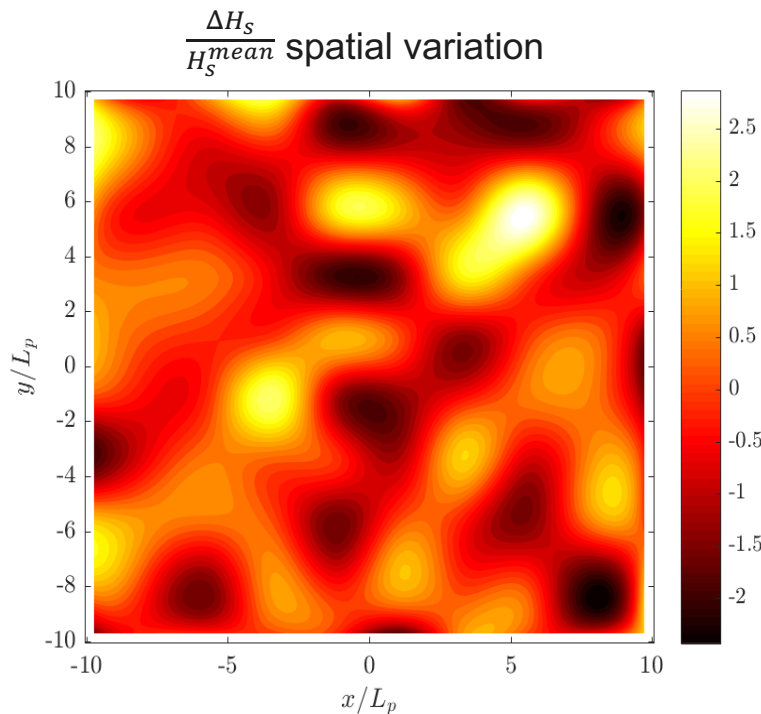


Methods of directional wave generation: DSM

Numerical Simulations with LRWT:

- Repeat time = 1024 s (3 hrs field equivalent) $\Delta f = 1/1024$ Hz
- Variations in $H_s \sim 5\%$ \longrightarrow unwanted in model testing
- Increase frequency discretization $\frac{1}{P\Delta f}$, $P = 2, 4, 8 \dots 48$

\longrightarrow practical problems with wave-makers



Methods of directional wave generation: SSM,RDM

- Single Summation Method (SSM):
 - Widely applied
 - Division in frequency bands,
 - Each frequency – 1 direction
 - No problems with ergodicity
 - Sensitive to the discretization of the frequency spectrum

$$\eta(x, y, t) = \sum_{i=1}^N A_i \cos[\omega_i t - k(x \cos \theta_i + y \sin \theta_i) + \epsilon_i] \quad (4) \text{ where}$$

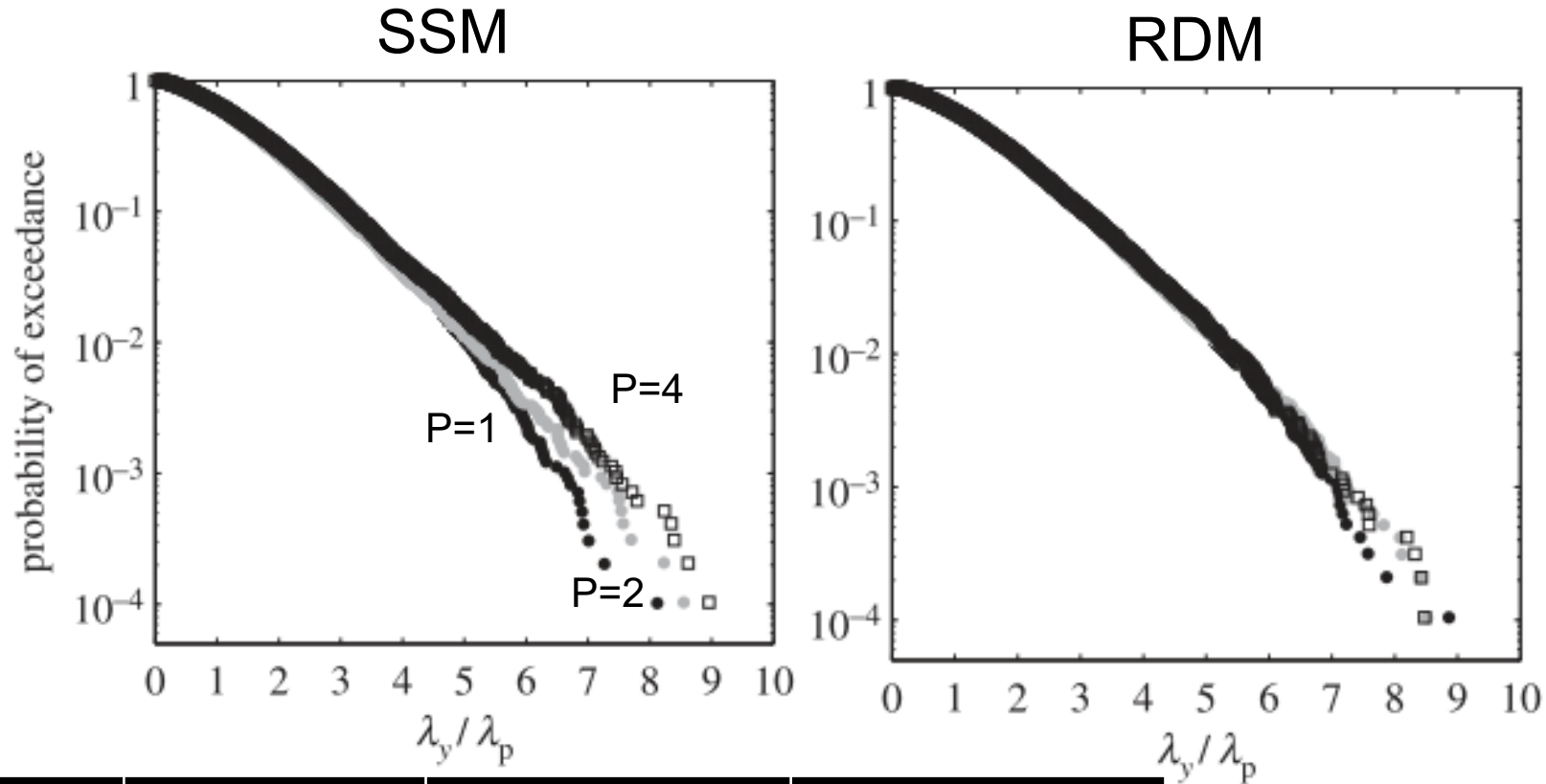
$\omega_i = \frac{i \Delta \Omega}{M}$, where $\Delta \Omega$ is the width of the band.

- Random Directional method (RDM)
 - Same as above, single summation
 - Each frequency – 1 direction across the whole spectrum
 - Directions chosen randomly from a weighting function
 - Less sensitive to discretization



Methods of directional wave generation: SSM, RDM

Normalised crest lengths



P	Δf (Hz)	no. directional components	no. frequency bands
1	1/1024	25	64
2	1/2048	30	108
4	1/4096	40	161

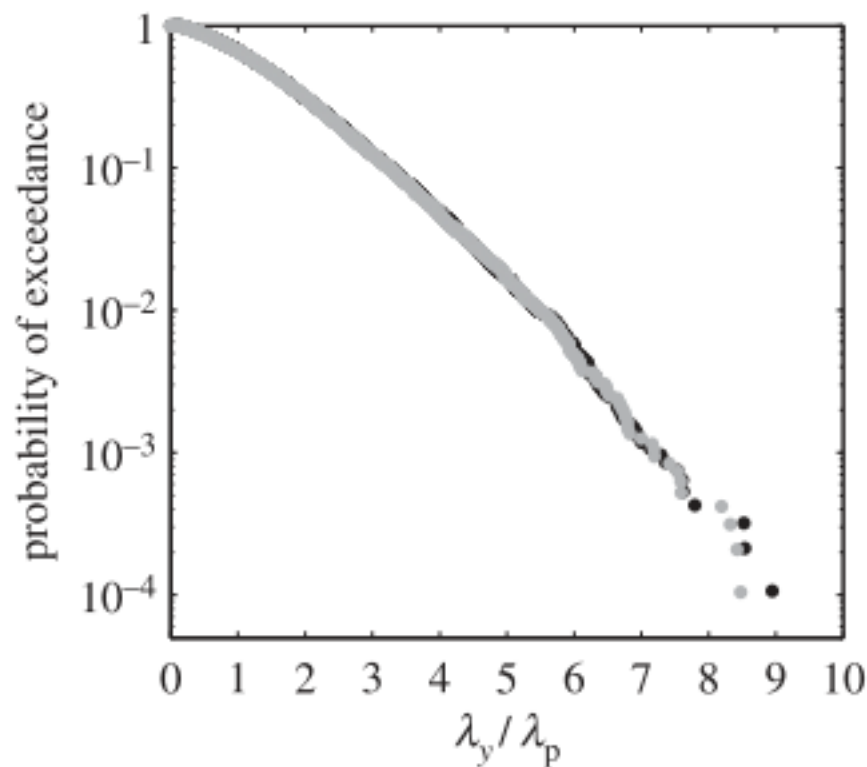


Methods of directional wave generation: SSM,RDM

Same results for finer discretization

SSM: $\Delta f = 1/8192$ RDM: $\Delta f = 1/4096$

SSM/RDM

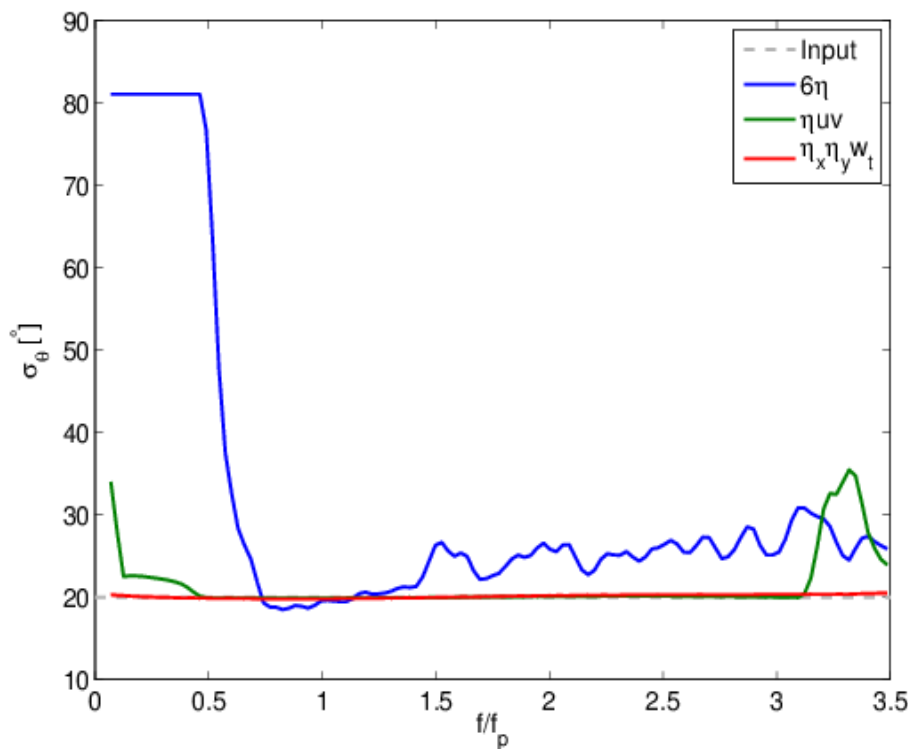
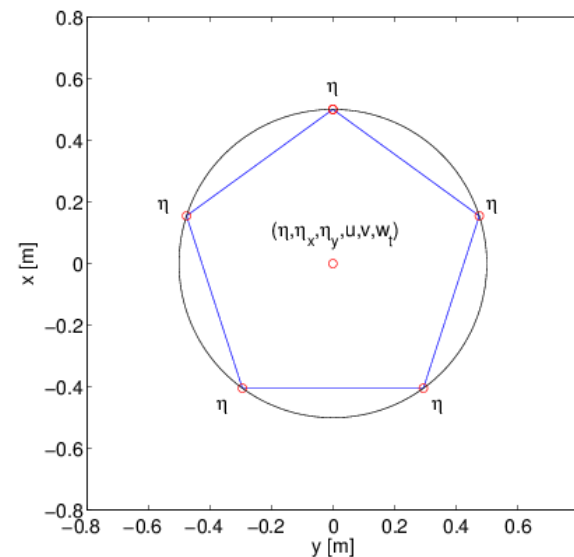


All methods give
same results for:

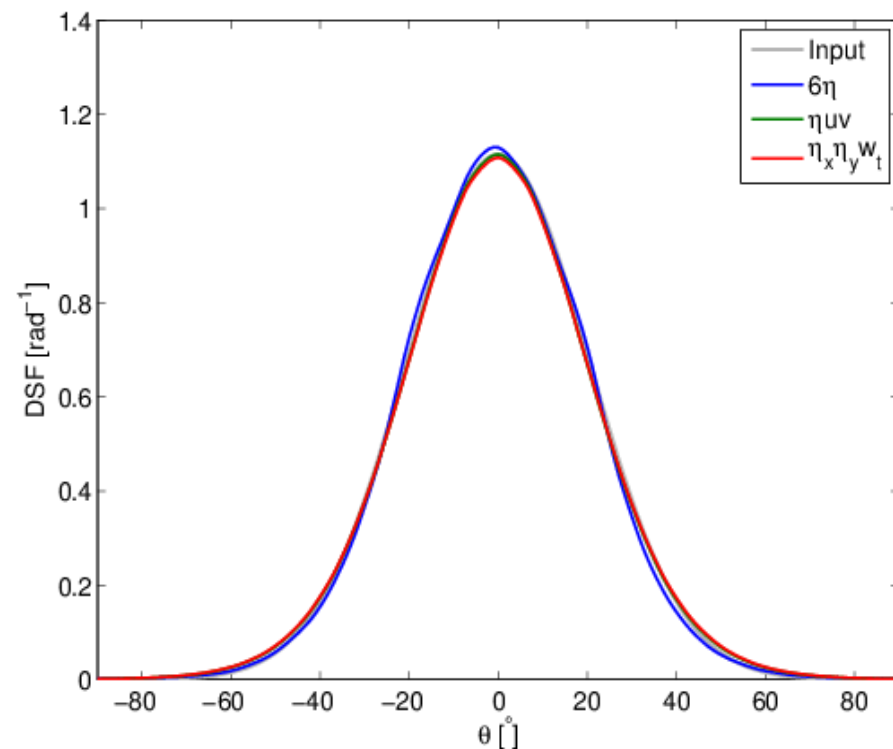
$$\frac{1}{\Delta f} \rightarrow 0$$

Directional analysis: input data

- $\sigma_\theta = 20^\circ$, 20 seeds x 1024 s
- calculations based upon the EMEP
- comparisons between various input data



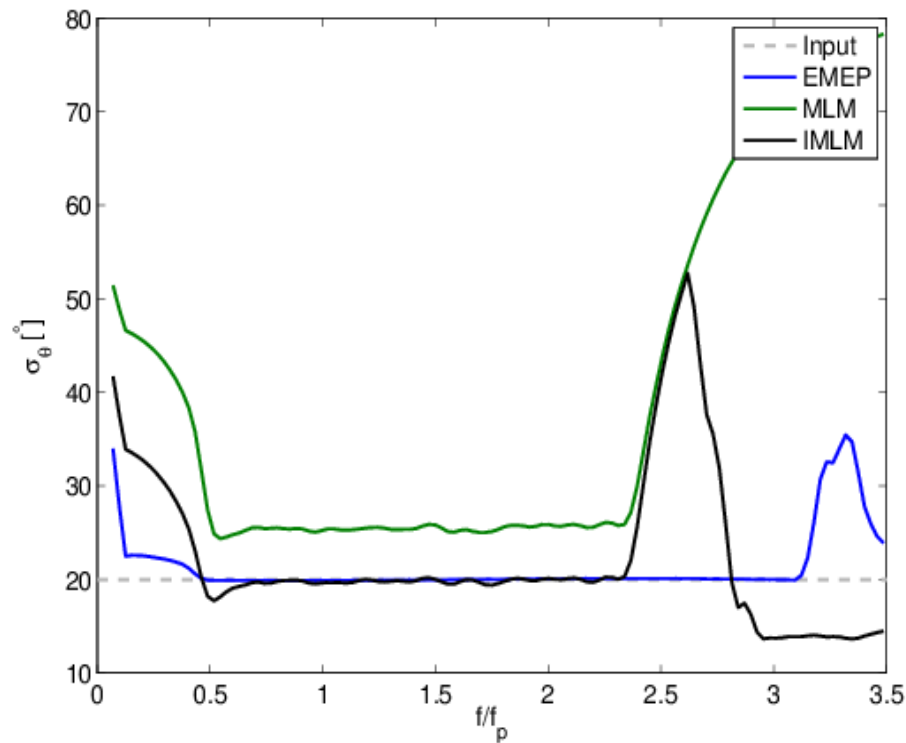
σ_θ vs. f/f_p



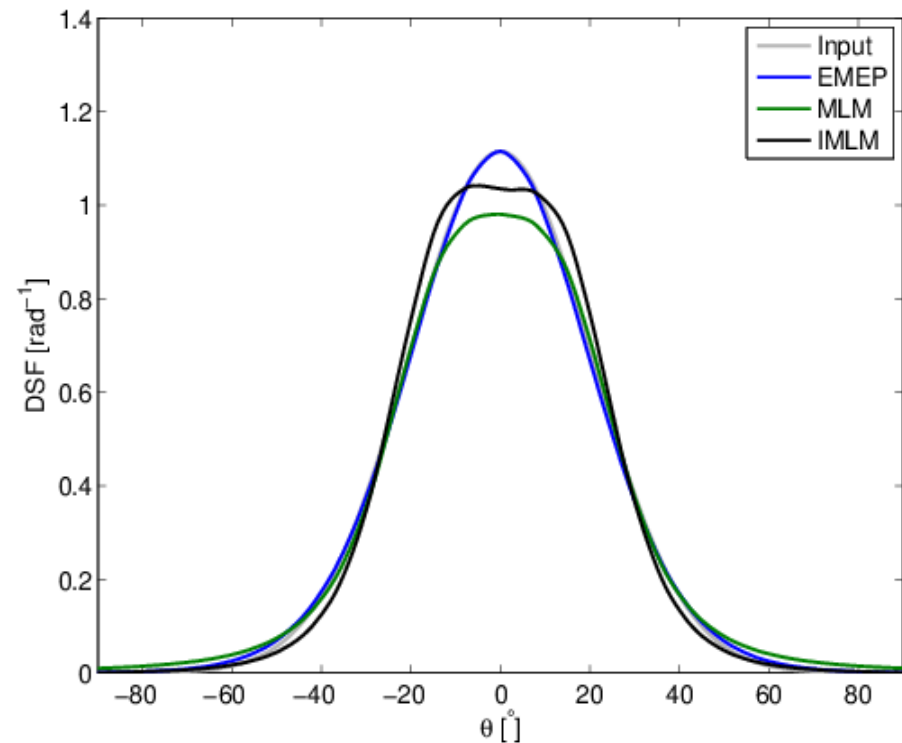
Directional spreading function, DSF

Directional analysis: methods

- $\sigma_\theta = 20^\circ$, 20 seeds x 1024 s
- calculations based upon η, u, v
- comparisons between various analysis methods



σ_θ vs. f/f_p

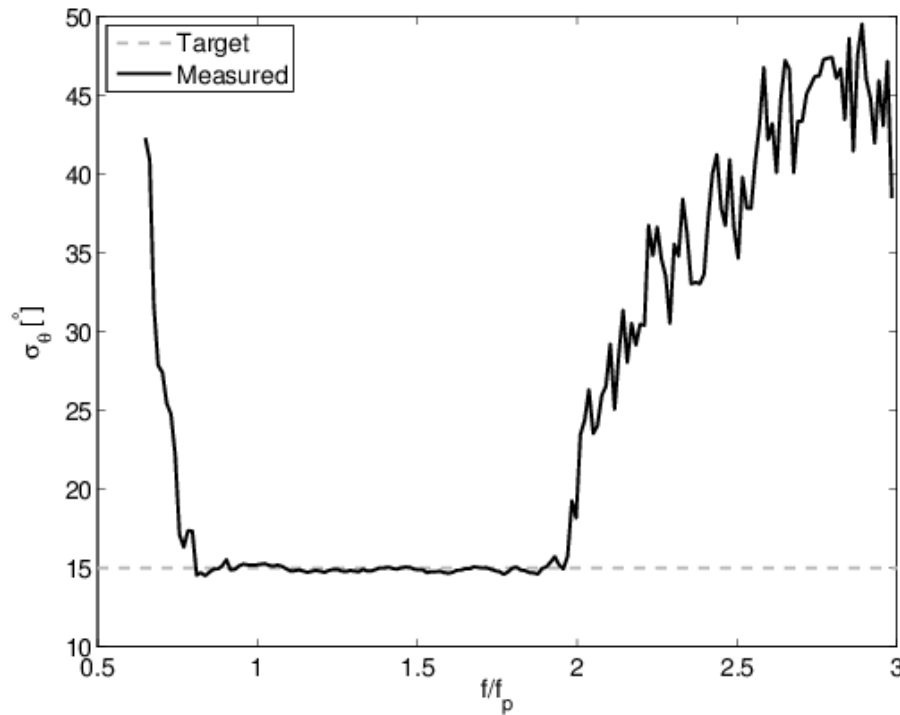


Directional spreading function, DSF

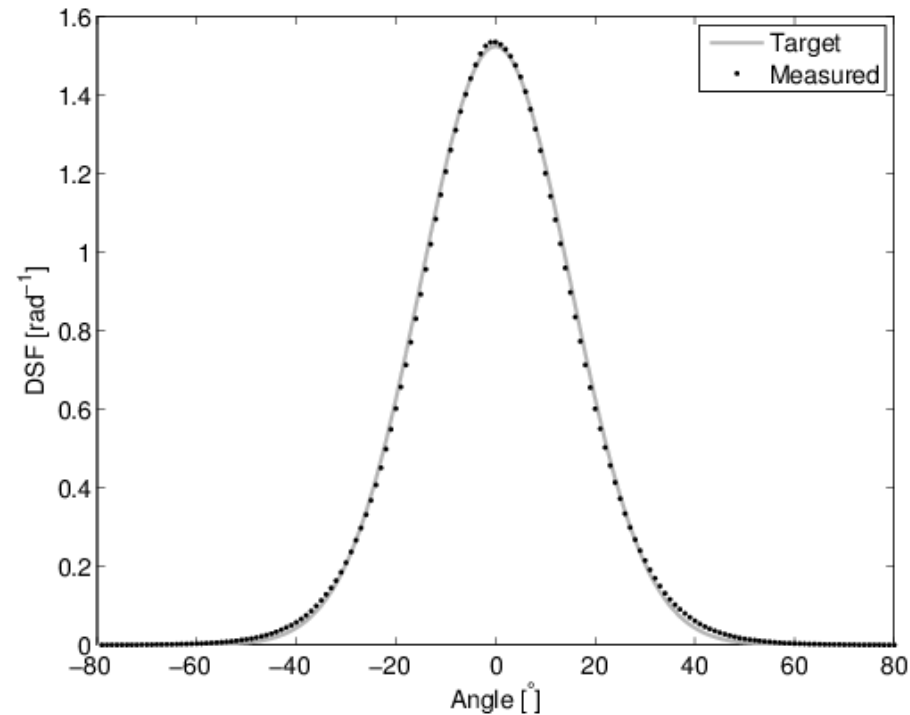
Directional analysis: sea-state

Comparisons to laboratory data ($H_s=10\text{m}$, $\frac{1}{2}H_s k_p=0.081$)

- $\sigma_\theta=15^\circ$
- input data: η, u, v
- calculated using the EMEP
- sea state generated using RDM



σ_θ vs. f/f_p

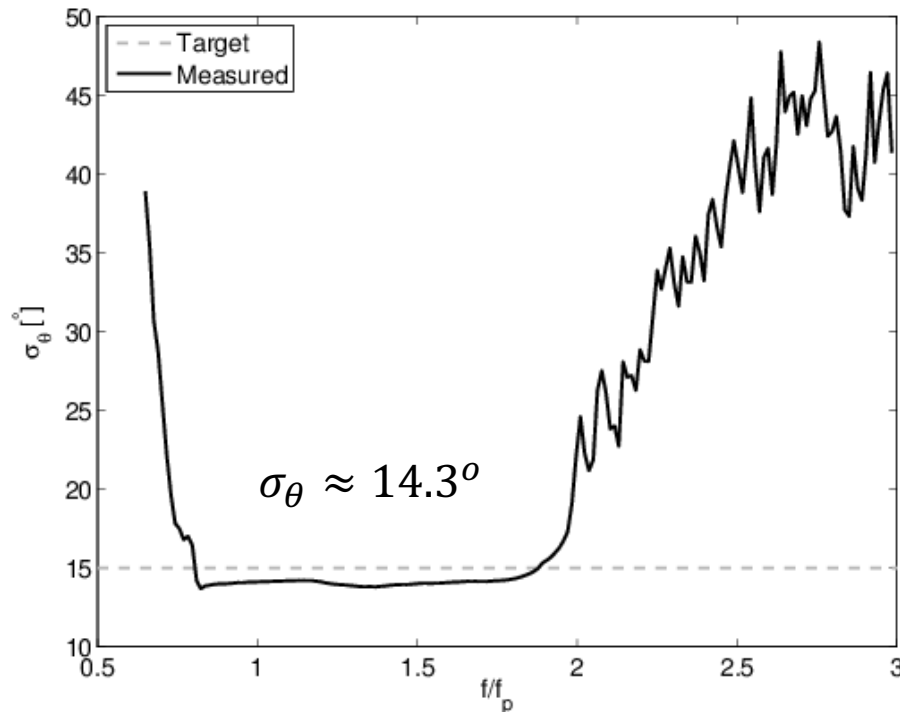


Directional spreading function, DSF

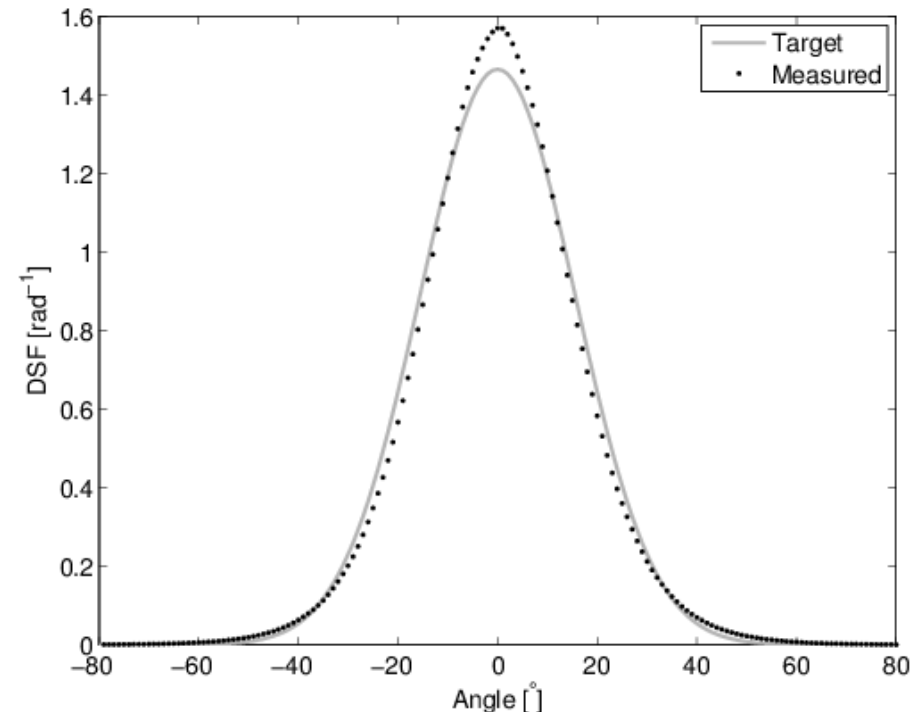
Directional analysis: sea-state

Comparisons to laboratory data ($H_s=15.0\text{m}$, $\frac{1}{2}H_s k_p=0.122$)

- $\sigma_\theta=15^\circ$
- input data: η, u, v
- calculated using the EMEP
- sea state generated using RDM



σ_θ vs. f/f_p

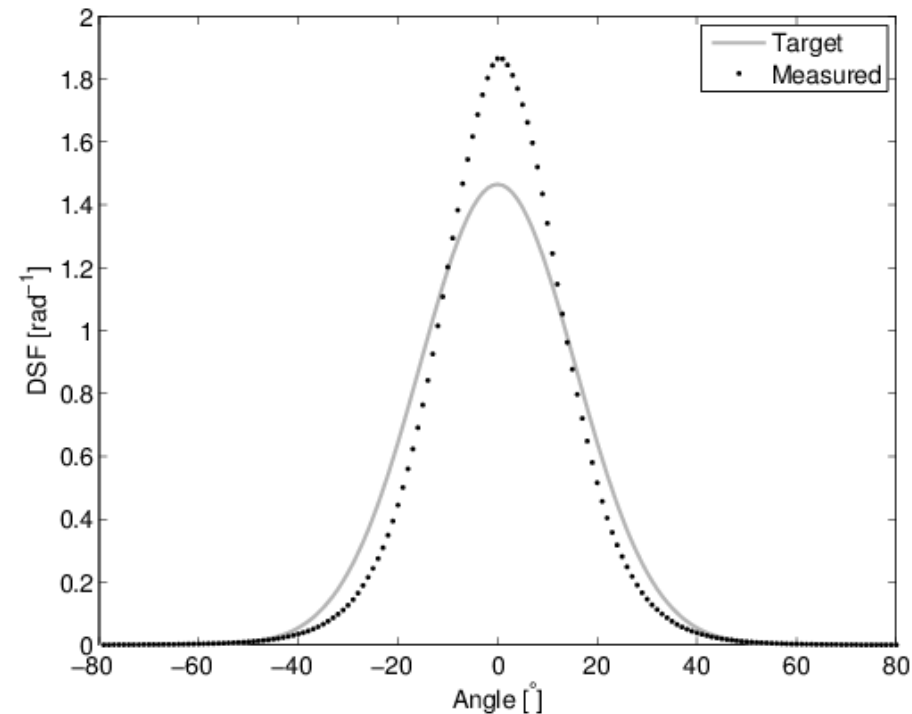
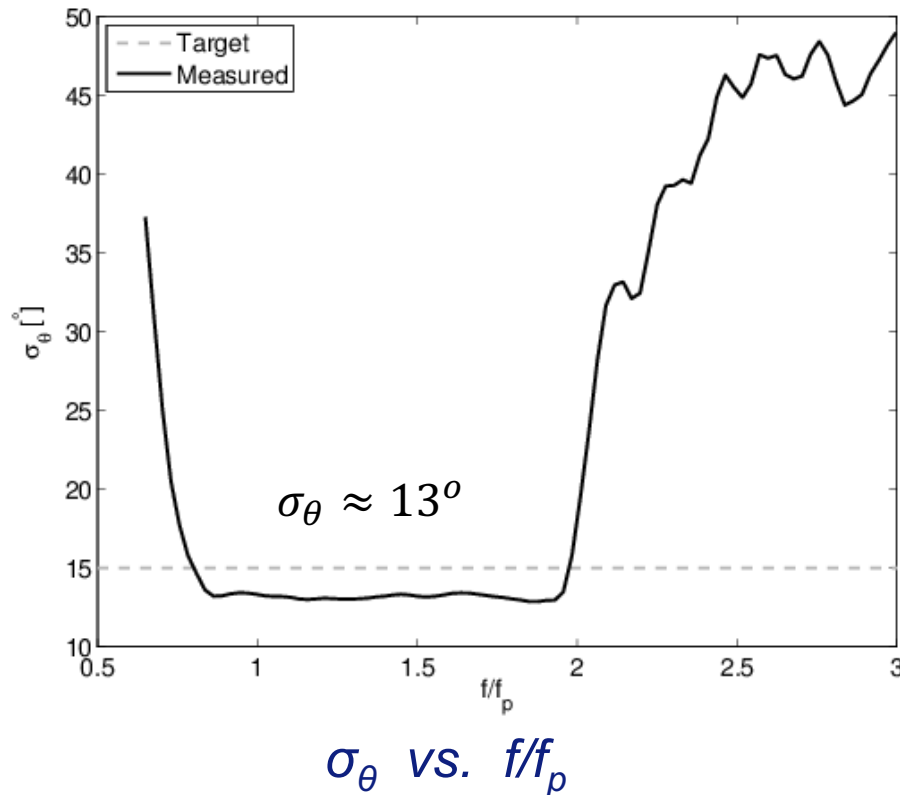


Directional spreading function, DSF

Directional analysis: sea-state

Comparisons to laboratory data ($H_s=20.0\text{m}$, $\frac{1}{2}H_s k_p=0.163$)

- $\sigma_\theta=15^\circ$
- input data: η, u, v
- calculated using the EMEP
- sea state generated using RDM

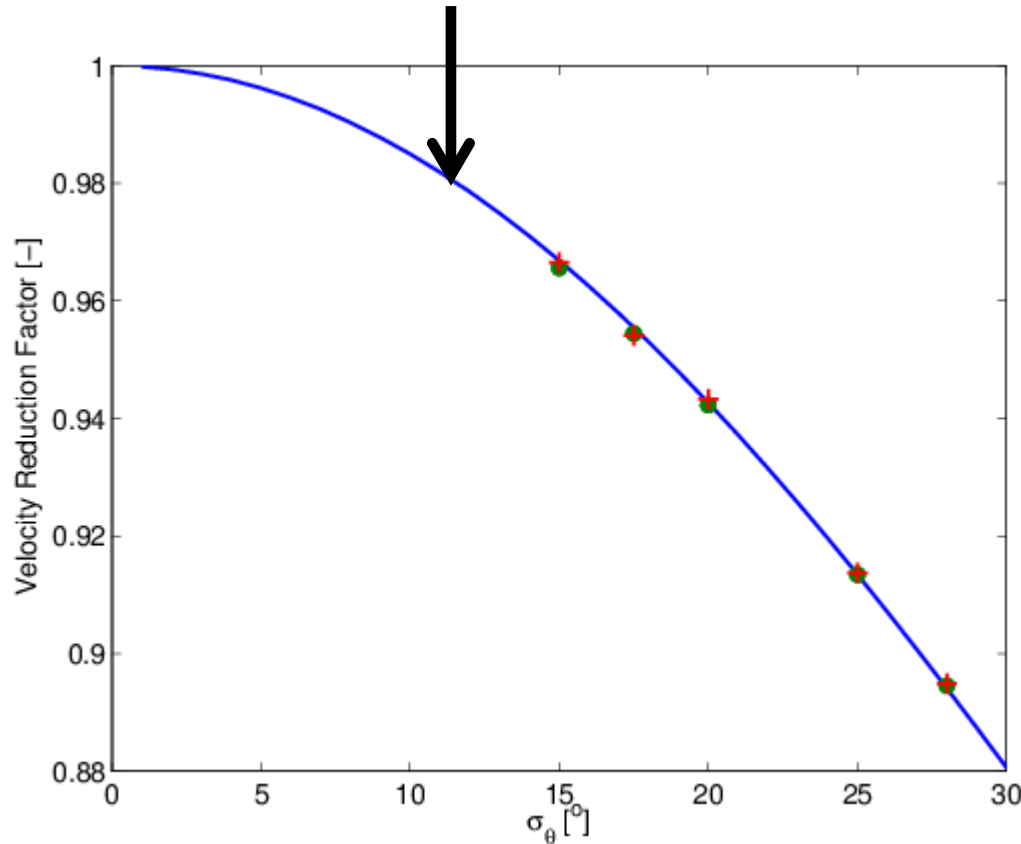


Directional analysis: alternative method

Based upon the velocity reduction factor (VRF)

$$\text{VRF} = \left[\frac{1}{2} + \frac{1}{2} \exp(-2\sigma_\theta^2) \right]^{\frac{1}{2}} \quad (5) \quad \text{or} \quad \text{VRF} = \frac{(\text{rms } u_x \text{ in directional sea})}{(\text{rms } u_x \text{ in uni-directional sea})}$$

(Tucker & Pitt, 2001)

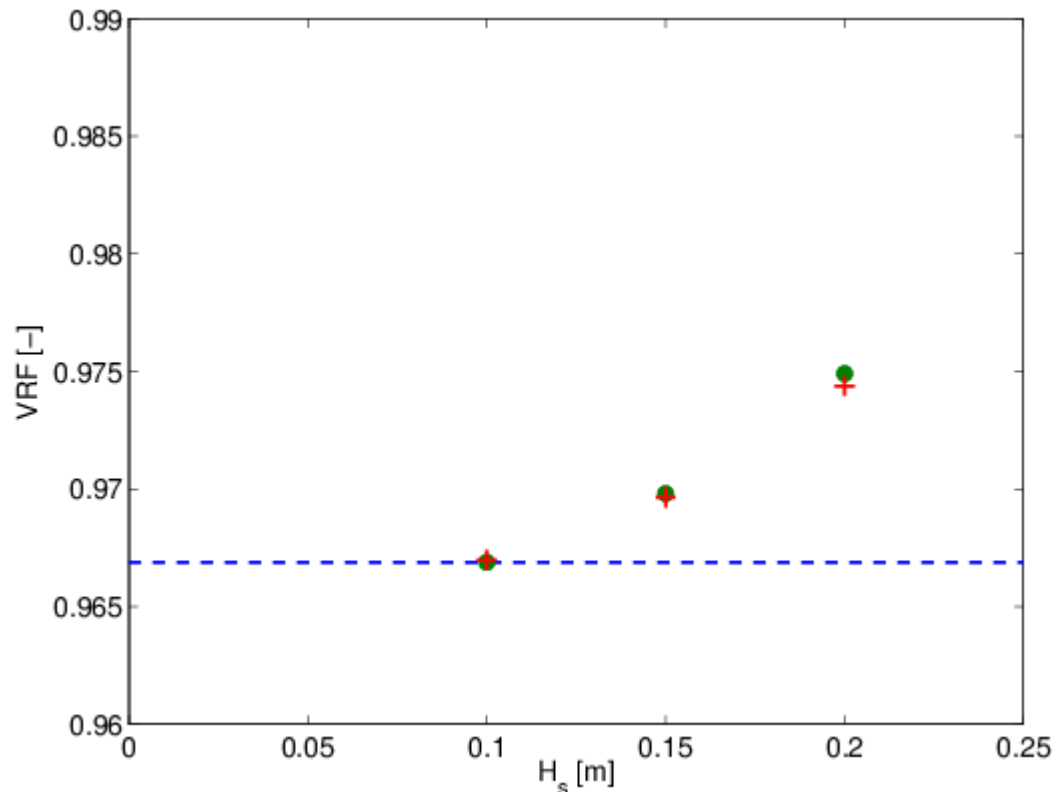


Based upon:

- the analysis of an entire wave record
- + the average of a wave-by-wave analysis

Directional analysis: alternative method

- comparisons to laboratory data
- VRF averaged over 20 x 3-hour seeds for each sea state
- changes with H_s

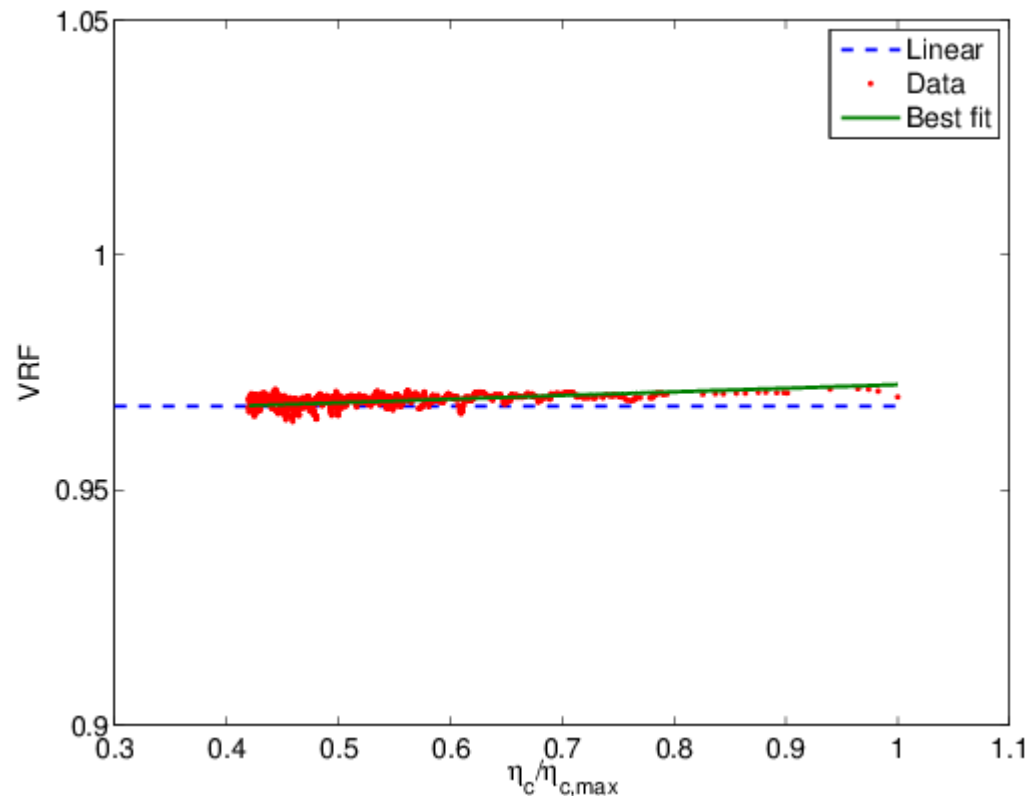


- - - linear input σ_θ
- velocity ratio
- + earlier EMEP σ_θ

Steeper sea-states are
more uni-directional

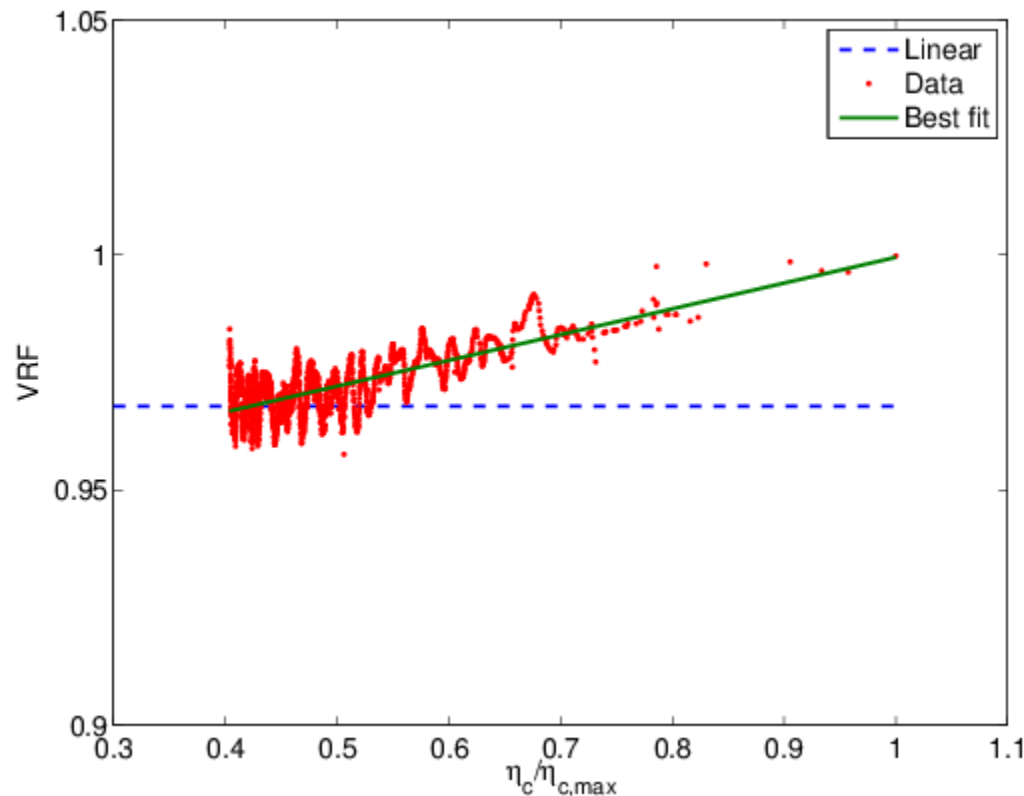
Velocity reduction factor (VRF)

- Comparisons to laboratory data ($H_s=10\text{m}$, $\sigma_\theta=15^\circ$, $\frac{1}{2}H_s k_p=0.081$)
- VRF calculated for individual waves
- Plotted in terms of the normalised crest elevation, $\eta_c/\eta_{c,max}$



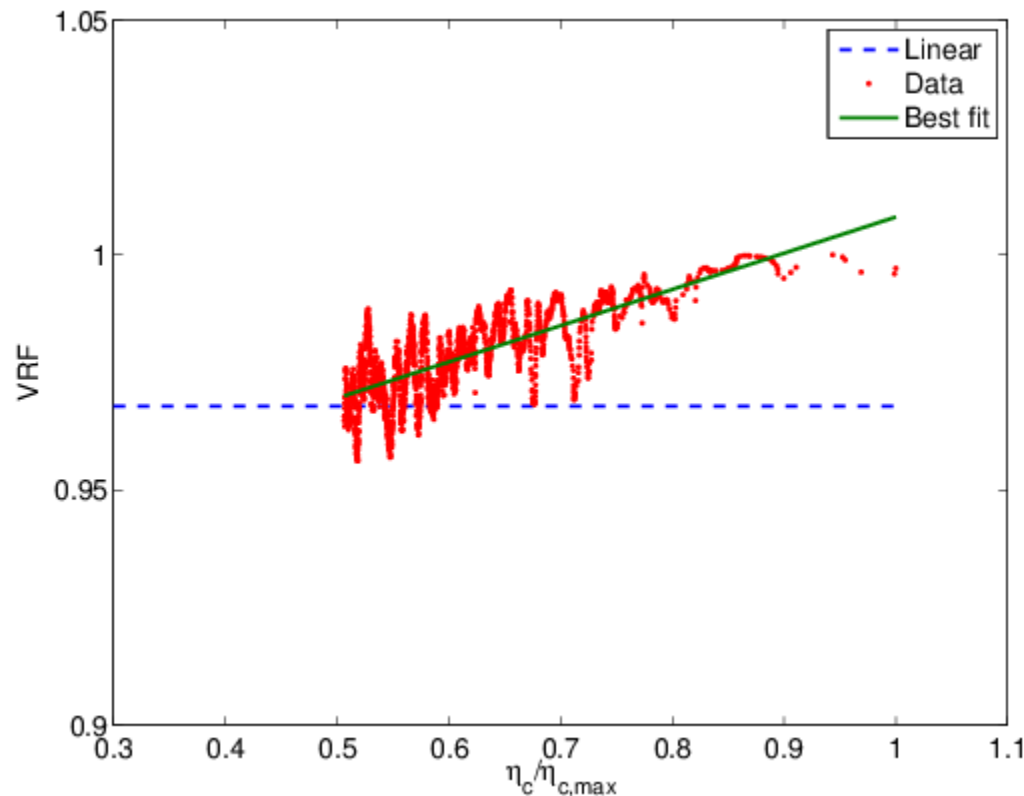
Velocity reduction factor (VRF)

- Comparisons to laboratory data ($H_s=15\text{m}$, $\sigma_\theta=15^\circ$, $\frac{1}{2}H_s k_p=0.122$)
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Velocity reduction factor (VRF)

- Comparisons to laboratory data ($H_s=20\text{m}$, $\sigma_\theta=15^\circ$, $\frac{1}{2}H_s k_p=0.163$)
- VRF calculated for individual waves
- Plotted in terms of the normalised crest elevation, $\eta_c/\eta_{c,max}$



Concluding remarks

- Preferred method of directional wave generation: RDM
 - Computationally efficient
 - Ergodic
 - Easy to implement
- Directional spreading decreases in steeper sea-states
- Large individual waves are less directionally spread
- Results agree with experimental and numerical studies (e.g. Johannessen & Swan, 2001 & 2003; Adcock, et al. 2012 & 2015)
- Extension to intermediate and shallow water depths through LoWiSh JIP (currently restricted)
- Part of results available in Proc. Royal Society:

"A laboratory study of nonlinear changes in the directionality of extreme seas" (2017), M.Latheef, C.Swan, J.Spinneken



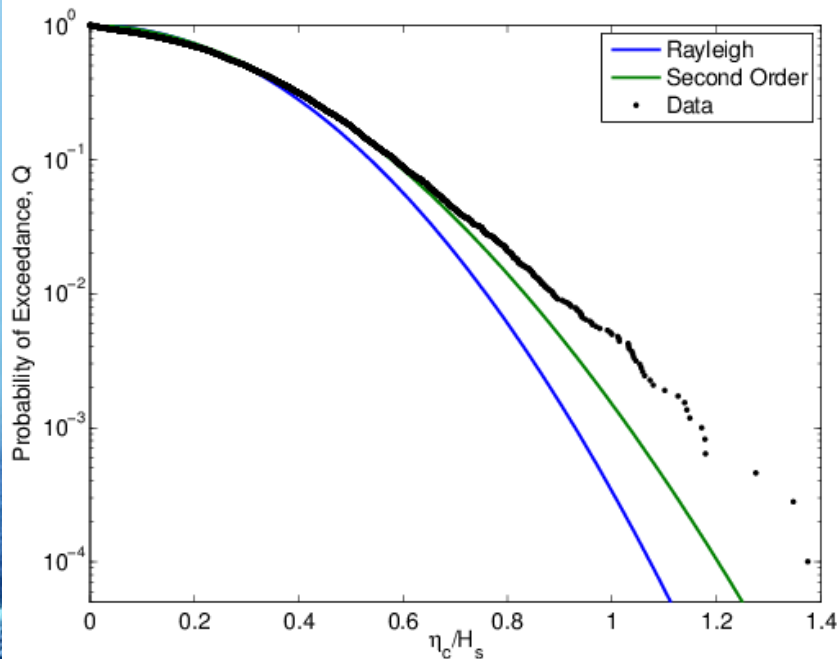
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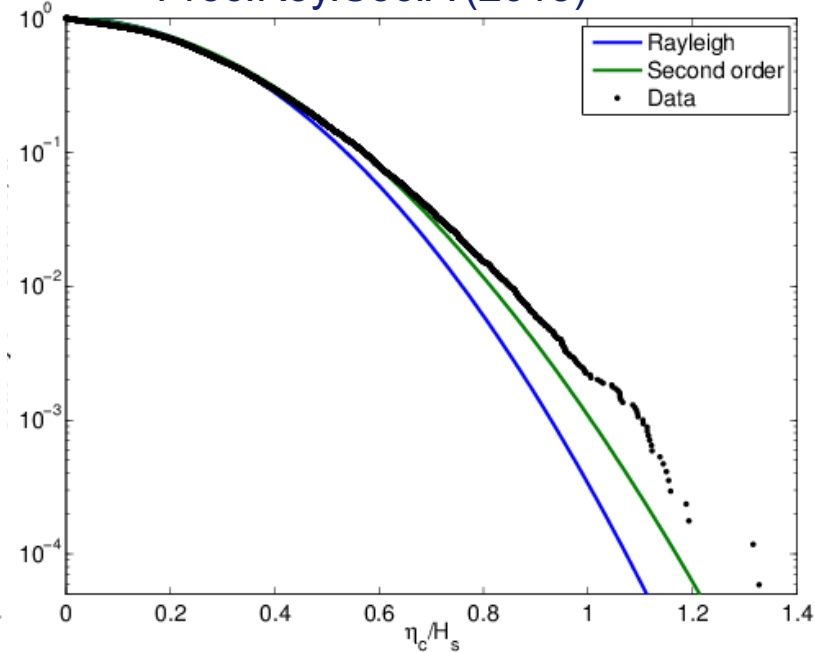
Short-term distribution of crest heights

- Effects beyond $O(a^2k^2)$
- Both in field data (North Sea) and laboratory data (ICL)

Field data ($H_s > 12$ m)
Christou & Ewans (2012)
CREST JIP

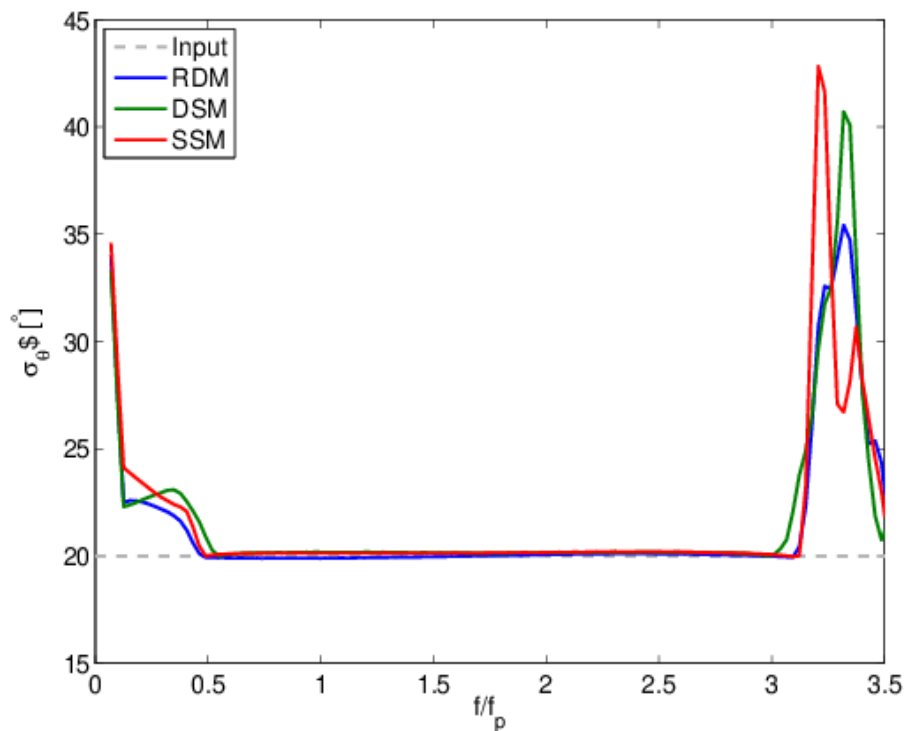


Experimental data ($H_s = 12.5$ m,
 $T_p = 16$ s, $\sigma_\theta = 15^\circ$)
Latheef & Swan
Proc.Roy.Soc.A (2013)

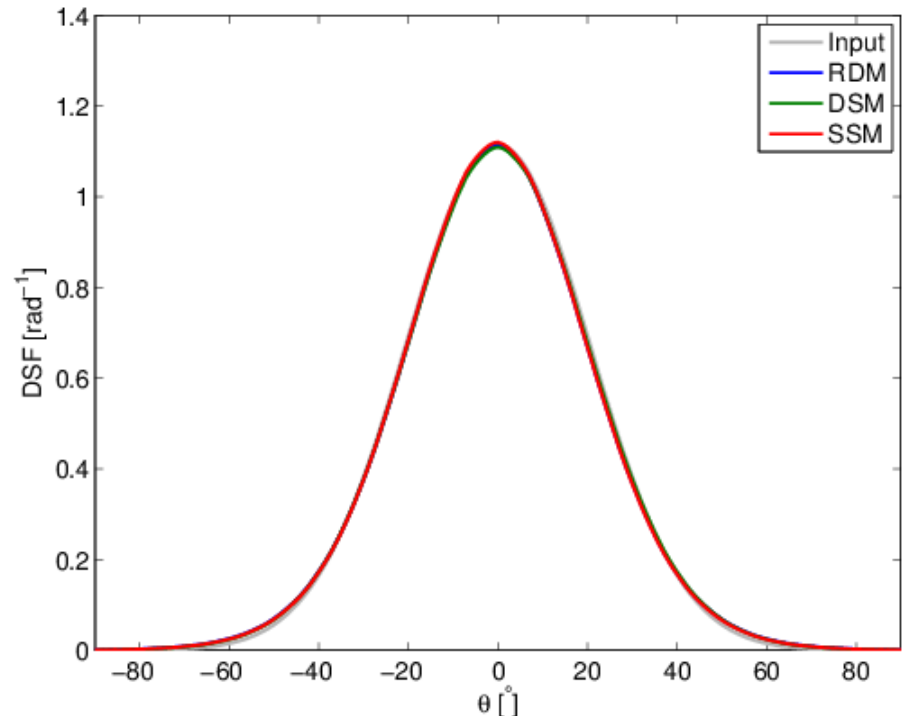


Directional analysis: Generation method

- $\sigma_\theta = 20^\circ$
- calculations based upon the EMEP
- comparisons between different methods of directional simulation



σ_θ vs. f/f_p



Directional spreading function, DSF

Directional spectrum

Given the frequency spectrum, $S_{\eta\eta}(\omega)$, the directional spectrum is:

$$F(\omega, \theta) = S_{\eta\eta}(\omega)D(\omega, \theta),$$

where $D(\omega, \theta)$ is the directional spreading function (DSF).

In terms of Fourier series:

$$D(\omega, \theta) = \frac{1}{\pi} \left\{ \frac{1}{2} + \sum_{n=1}^{\infty} A_n(\omega) \cos n\theta + B_n(\omega) \sin n\theta \right\}, \quad (1)$$

where $A_n(\omega) = \int_{-\pi}^{\pi} D(\omega, \theta) \cos n\theta$ and $B_n(\omega) = \int_{-\pi}^{\pi} D(\omega, \theta) \sin n\theta$

In this study: Gaussian DSF - frequency independent

$$D(\omega, \theta) = \frac{A}{\sigma_{\theta} \sqrt{2\pi}} e^{\left[-\frac{(\theta - \theta_m)^2}{2\sigma_{\theta}^2} \right]} \quad (2)$$

with σ_{θ} the standard deviation and θ_m the mean wave direction

RMS spreading: $\sigma_{\theta} = \sigma_1(\omega) = \sqrt{2[1 - \sqrt{A_1^2(\omega) + B_1^2(\omega)}]}$



Earlier work:

- Numerical calculations of focused waves (spectral model - BST)
- Local reduction in directional spreading
- Supporting laboratory data (Johannessen & Swan, 2001 & 2003)

