Role of Coastal HF Radars in Storm-Surge Forecasting

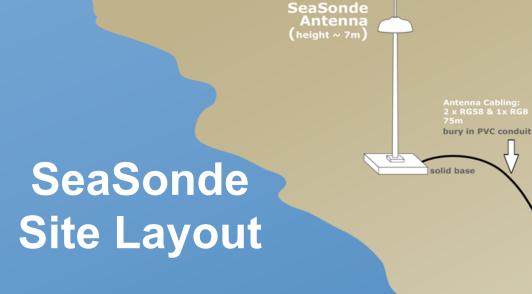
Dr. Don Barrick – President, CODAR Ocean Sensors Waves & Storm Surge Workshop -- November 10, 2015

• Where are Radars Useful in Three Phases of Storm Surge?

- Inundation or rise: the most important need to know how high water will rise in order to manage impacts
- Steady-state or equilibrium max wind speed reached, bottom return flow balances surface onshore flow
- Relaxation wind drops, shoreward surface flow ceases, gravity allows all flow offshore
- Where Can Radars Lead to Useful Short-Term Forecasts?



shed/ building



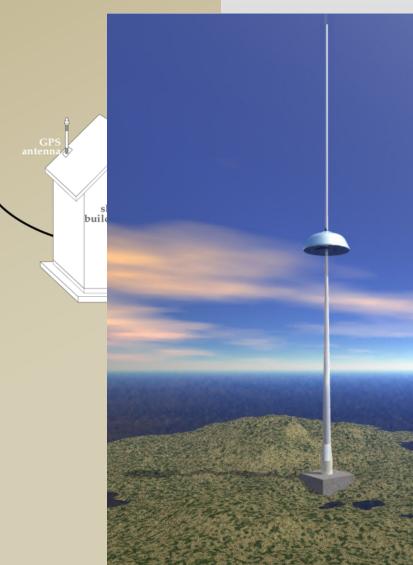
Climate-Controlled Shelter with Power, Communications

> 11 MHz: Single T/R antenna









A hurricane-hardened "NOAA Sentinel" water-level observing station.



140 Candidate SeaSonde HF Radars for Near-Field Storm-Surge Modeling & Forecasting

Canada

United States

Mexico

North acific)cean "I am a SeaSonde coastal HF radar. What do I see and how can I help with short-term storm-surge forecasts?"



- Surface velocities mapped in bands out to ~200 km every hour
- Thence, deduce onshore water transport in upper layer forced by winds and waves
- Isn't this all that's needed for the space/time domain of HFR during 'inundation phase'?
- If so, winds/waves unnecessary for short-term forecasts (<12 h)

Example for Hokkaido of Bands and Fitting of

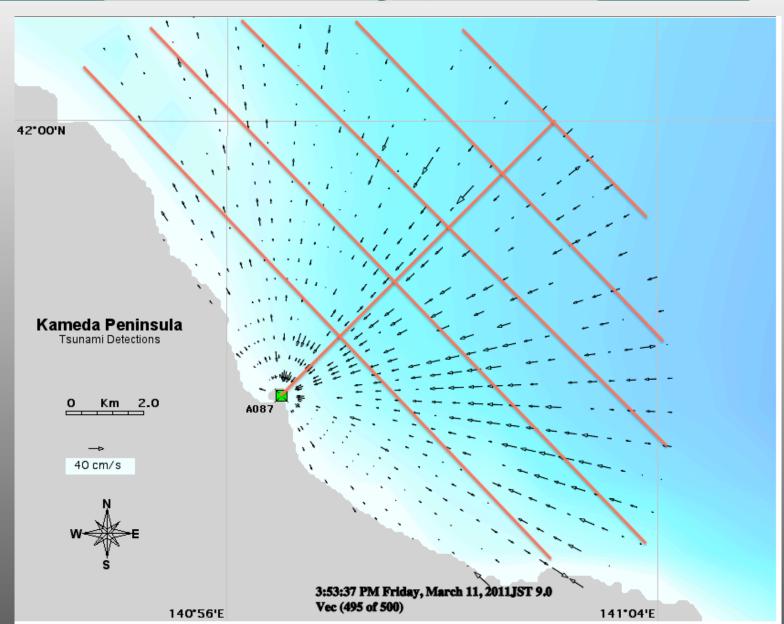
Radials from Single SeaSondes

EAN

0 C

ODAR

SENSORS



Inundation Phase-1: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Conservation of Momentum (Newton's Second Law) [Resio et al., 2008]

$$\left(\frac{\partial u(x,z,t)}{\partial t} = \frac{1}{\rho_s} \frac{\partial \tau_x(x,z,t)}{\partial z}\right)$$

Acceleration Related To Wind Stress

Inundation Phase-1: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Conservation of Momentum (Newton's Second Law) [Resio et al., 2008]

• Reynolds Stress Related to Eddy Viscosity, *µ*

(2)
$$\tau_x(x,z,t) = \mu \frac{\partial u(x,z,t)}{\partial z} \qquad 10^{-5} \le \mu \le 10^{-3} \quad \frac{\mathrm{m}^2}{\mathrm{s}}$$

Inundation Phase-1: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Conservation of Momentum (Newton's Second Law) [Resio et al., 2008]

. Republic Stress Related to Riddy Miscosity, μ

• Combine to Get Parabolic Partial Differential Equation (Heat Eq.)

$$\left(\frac{\partial u(x,z,t)}{\partial t} = \mu \frac{\partial^2 u(x,z,t)}{\partial z^2}\right)$$

(3)

Horizontal x-Variation Is Slow

Inundation Phase-2: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Conservation of Water (incompressibility) [Resio et al., 2008]

$$\frac{\partial \eta(x,t)}{\partial t} = -\int_{-\eta}^{d} \frac{\partial u(x,z,t)}{\partial x} dz$$

When you squeeze water at sides of column, sea level must rise

• Surface (storm-surge) height: $\eta(x,t)$

Inundation Phase-2: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Conservation of Water (incompressibility) [Resio et al., 2008]

- Surface (storm-surge) height: $\eta(x,t)$
- Assume a plausible model for velocity:

$$u(x,z,t) = a \exp\left(\frac{z}{\delta(x,t)}\right) \tanh\left(\frac{\tau_o(x,t)t}{a\rho_s\delta(x,t)}\right) \text{ for } \begin{cases} x>0\\ z<0\\ t>0 \end{cases}$$

Inundation Phase-2: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Conservation of Water (incompressibility) [Resio et al., 2008]

- Surface (storm-surge) height: $\eta(x,t)$

 Typically water final velocity at surface due to wind stress is approximately: a ~ 0.03V₁₀

Inundation Phase-3: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Integrate model (5) within (4) over z (positive downward) to get at upper-layer e-folding depth at surface, $\delta(x,t)$

$$\frac{\partial \eta(x,t)}{\partial t} = -\delta(x,t)\frac{\partial u(x,0,t)}{\partial x}$$

Inundation Phase-3: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Integrate model (5) within (4) over *z* (positive downward) to get at upper-layer e-folding depth at surface, $\delta(x,t)$

• Differentiate model (5) PDE (Heat Eq.) (5) over z:

()
$$\frac{\partial u(x,0,t)}{\partial t} = \frac{\mu}{\delta^2(x,t)}u(x,0,t)$$

Inundation Phase-3: Transport Leading to Surge-Level Rise Occurs in Upper Layer

• Integrate model (5) within (4) over *z* (positive downward) to get at upper-layer e-folding depth at surface, $\delta(x,t)$

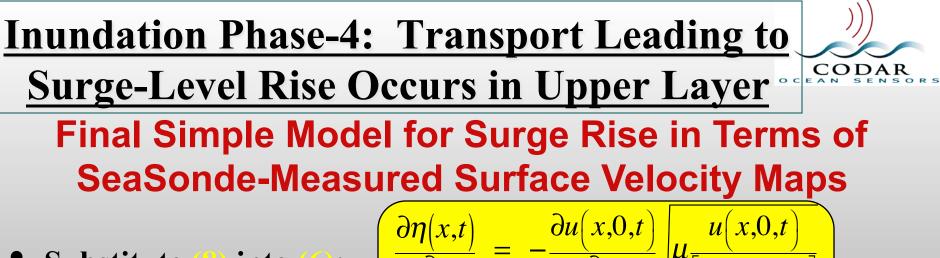
Differentiate model (5) FDE (Heat Eq.) (5) over **z**:

$$\frac{\partial u(x,0,t)}{\partial t} = \frac{\mu}{\delta^2(x,t)} u(x,0,t)$$

• Solve the above equation to get e-folding upper-layer depth:

$$\delta^{2}(x,t) = \mu \frac{u(x,0,t)}{\left[\frac{\partial u(x,0,t)}{\partial t}\right]}$$

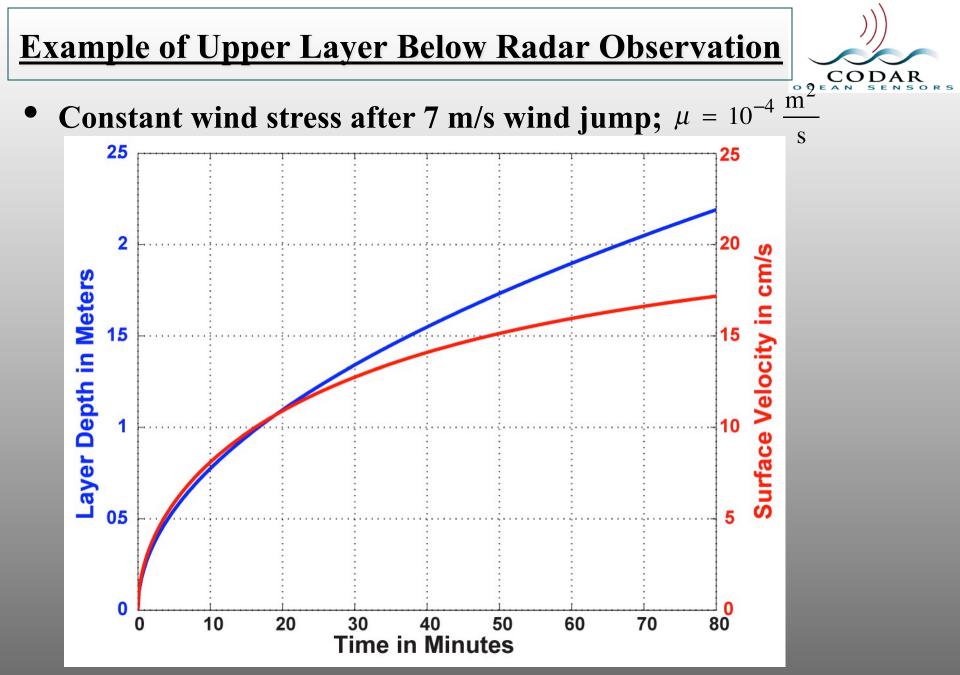
(8)



• Substitute (8) into (6):

 $\frac{\partial \eta(x,t)}{\partial t} = -\frac{\partial u(x,0,t)}{\partial x} \sqrt{\frac{u(x,0,t)}{\left[\frac{\partial u(x,0,t)}{\partial t}\right]}}$

- Radar measures: $\mu(x_i, 0, t_j)$
- Integrate above equation over time to get: $\eta(x_i, t_j)$
- How far into future time can this be extended based on past measurement?



For later times Coriolis/Ekman effect becomes important

Conclusions and Way Forward



- >100 U.S. coastal SeaSonde HF radars observe onshore surface flow/transport out to 200 km, continuous over space & time
- <u>Premise</u>: these can be used to forecast sea level rise during inundation phase, perhaps out to 12 24 hours
- We derived simple hydrodynamic example to demonstrate this
- Is this the final/best solution? Absolutely not, it is a start!
 - This is a "first-order" attempt
 - Better but still simple numerical model will include higher-order effects, allow longer forecasts
- Supplement Larger-Scale Wind/Wave Storm Surge Models with Local Radar-Forecasted Near Fields