CBLAST-Hurricane and High-Resolution Fully Coupled Atmosphere-Wave-Ocean Models

Shuyi S. Chen

Rosenstiel School of Marine and Atmospheric Science University of Miami (W.Zhao, M.Donelan, J.Price, E.Walsh, H.Tolman, &CBLAST team)

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In the eye of Katrina

Motivation:

> to better understand and predict hurricane structure and intensity

Hurricane Maximum Potential Intensity (MPI) Emanuel (1988), Holland (1997), etc. $MPI = f \{ (C_k/C_D), \epsilon, SST, RH \}$ RH = relative humiditySST = sea surface temperature $\epsilon = (T_B - T_O)/T_B \text{ (thermodynamic efficiency)}$ $C_k \text{ and } C_D \text{ are exchange coefficients of enthalpy and momentum fluxes}$

Why most hurricanes do not reach their MPI?

- Inner core (eye and eyewall) dynamics
- Environmental conditions (atmosphere and ocean)

Coupled Atmosphere-Wave-Ocean Modeling System for Hurricane Predictions



Summary of Conclusions

- Fully coupled, high-res models improve intensity prediction and provide a tool to better understand physical processes that lead to extreme winds and heavy rain in hurricanes
- Wind-wave coupling with 2D wave stress is critical in forecasting hurricane surface winds
- C_k/C_D varies spatially in hurricanes from 0.4-0.8 (inner-outer)

"Ensemble" Model Forecasts of Storm Intensity During Hurricane Katrina



Impact of Model Grid Resolution on Hurricane Forecast



Airborne radar <u>observed</u> rain in Hurricane Floyd (1999)





High-resolution research model



Most global operational models



CBLAST-Hurricane

Coupled Atmosphere-Wave-Ocean Modeling



A goal of CBLAST is to better understand how hurricanes interact with the ocean, and to use this to improve hurricane forecast models.

➢ Through CBLAST we have improved our knowledge about the processes which fuel hurricanes (heat from the ocean), as well as the frictional forces (drag on the sea surface) which mix the ocean and result in extreme ocean waves.

Specific Objectives:

- Wind-Wave Coupling
- Effects of Sea Spray
- Atmosphere-Ocean Coupling

CBLAST (Coupled Boundary Layer Air-Sea Transfer)

100.

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Coupled Modeling System

• MM5 (PSU/NCAR)

(vortex-following nests with 45, 15, 5, and 1.67 km grid spacing, NCEP analysis and AVHRR or TMI/AMSR-E SST)

• WAVEWATCH III (NOAA/EMC)

(1/12°, 25 frequency bands, 48 directional frequency bands)

- HYCOM (UMiami/NRL) (1/12°, 22 vertical levels with 4-6 in the ocean mixed layer)
- 3DPWP (Price's 3-D Upper Ocean Circulation Models) (5 km, 25 vertical levels)

Ocean surface waves in Hurricane Frances (2004)





Uncoupled Atmosphere Model

Charnock Relationship: $z_0 = \alpha u_*/g$

Coupled Atmosphere-Wave Model

• Roughness Length (non-directional)

 $\tau = \tau_t + \tau_w \longrightarrow z_o$ (e.g., Janssen at ECMWF)

 $\mathbf{Z}_{\mathbf{0}}$ - wave-age dependent

• Stress Vector (directional)

$$\mathbf{M}_{\mathbf{x}} = - \tau_{\mathbf{x}}$$
$$\mathbf{M}_{\mathbf{y}} = - \tau_{\mathbf{y}}$$

 ${f au}_{f x}$, ${f au}_{f v}$ - components of stress from integral of

momentum input to the wave spectrum.

Wind-Wave Coupling Parameterization

Spectra Tail Parameterization:

$$\tau_{x} = g\rho_{w} \int_{0}^{\infty} \int_{-\pi}^{\pi} \frac{\gamma}{\omega} F(k, \vartheta) k_{x} k dk d\vartheta$$

X-component of stress from integral of momentum input to the spectrum:

$$\frac{\gamma}{\omega} = S \frac{\rho_a}{\rho_w} \left[\frac{U_{(\pi/k)} \cos \theta}{C(k)} - 1 \right] \cdot \left| \frac{U_{(\pi/k)} \cos \theta}{C(k)} - 1 \right]$$

Growth rate of each component from measurement of pressure-slope correlation (S-shelter coefficient, C-phase speed, U(π/k)-half wavelengh height wind speed)

$$F(k, \vartheta) = \alpha k^{-5} \sec h^2(\beta(\vartheta_k))$$

Spectrum of long waves from WAVEWATCH III; spectrum of short waves from fit to tail given below. α is adjusted to fit the highest modeled wavenumbers.

$$\beta = \frac{1.2}{\cos^{-1}(C/U)}; C/U < 0.9$$

 β is the spreading function for the short waves.





stress across wind direction (%) 1200 UTC 31 AUG 04



Hurricane Frances (2004)

Frances Uncoupled $C_{D}(10^3)$ for 1200 UTC 31 Aug 2004



Coupled Frances Ocean/Wave Coupled $C_{\rm D}(10^3)$ for 1200 IJTC 31 Aug 2004



Black et al (2007), Drennan et al (2007), French et al (2007)



Hurricane Frances (2004)



Hurricane Frances (2004)



Coupled Model Forecast of Hurricane Katrina



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Hurricane Frances (2004) Model Mode Oh Coupling Atmosphere and Ocean from Low Winds to Hurricane



BAMS issue on CBLAST:

Chen et al., 2007: The CBLAST-Hurricane Program and the next-generation fully coupled atmosphere-waveocean models for hurricane research and prediction. *BAMS*, *311-317*.

Black et al., 2007: Air-Sea Exchange in Hurricanes: Synthesis of Observations from the Coupled Boundary Layer Air-Sea Transfer Experiment, *BAMS*, 357-374.

Edson et al., 2007: The Coupled Boundary Layers and Air-Sea Transfer Experiment in Low Winds (CBLAST-LOW). *BAMS, 346-*356.

White cap in hurricane environment

14

12

10



Wave energy dissipation

, from WW3+Tail





Whitecap coverage vs. U₁₀



The fractional whitecap coverage found from our video images falls consistently below the curve fit by Wu of some data from Monahan, and Toba and Chaen.

It is consistent with video image processing performed on higher and lower wind speeds, and the MIZEX and Gulf of Alaska experiments.

Note that Wu's fit would give 100% coverage at approx. 35 m/s!

Conclusions and Challenges

- Fully coupled, high-res models improve intensity prediction and provide a tool to better understand physical processes that lead to extreme winds and heavy rain in hurricanes
- Wind-wave coupling with 2D wave stress is critical in forecasting hurricane surface winds
- C_k/C_D varies spatially in hurricanes from 0.4-0.8 (inner-outer)
- Observations at the air-sea interface for U₁₀ > 30 m/s
- Wave-breaking and sea-spray remains to be an unresolved issue and need to be further examined in the coupled model (e.g., link to wave dissipation and breaking parameters, etc.).
- Accurate surface flux (especially enthalpy flux) and PBL measurements to evaluate coupled model results and improve model parameterizations
- Lack of good model initial conditions we need robust and efficient data assimilation system for coupled models